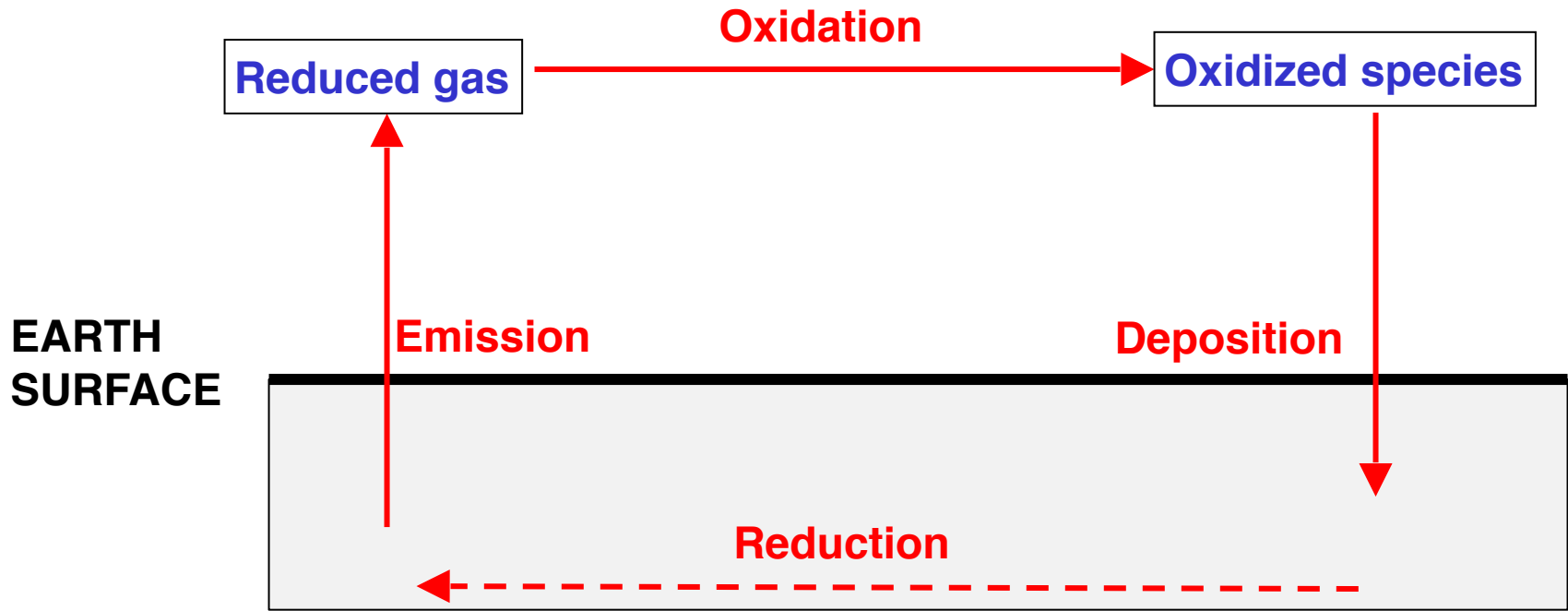
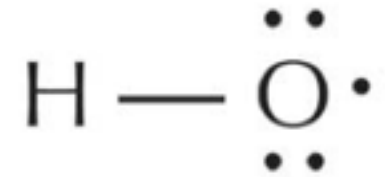


The troposphere as an oxidizing medium

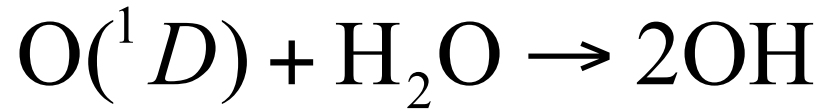
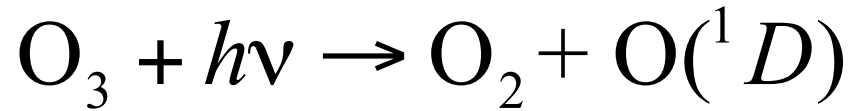


- Oxidation cleans the atmosphere of environmentally harmful gases:
 - Greenhouse gases such as methane, halocarbons
 - Toxic gases such as CO, benzene
 - Ozone-depleting gases such as hydrochlorofluorocarbons (HCFCs)
- Where the oxidation takes place affects patterns of deposition:
 - Acid rain
 - Nitrogen, mercury deposition

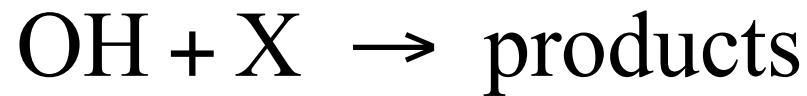
Main tropospheric oxidant: the OH radical



Source:



Sink:



where X is almost any non-radical reduced species

Lifetime:

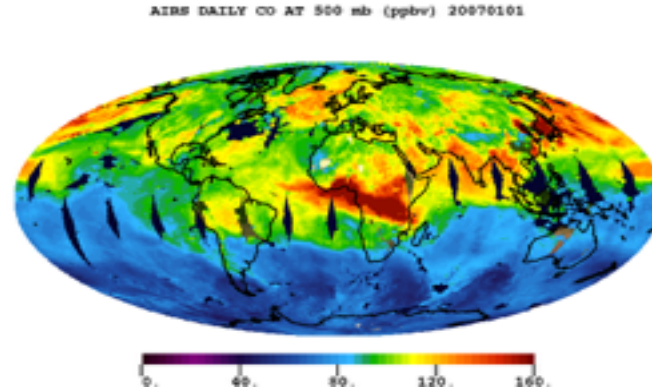
$$\tau_{\text{OH}} = \frac{1}{\sum_i k_i [\text{X}_i]} \sim 1 \text{ s} \quad \blackrightarrow \text{ concentrations of OH are low, variable}$$

Principal OH sinks:

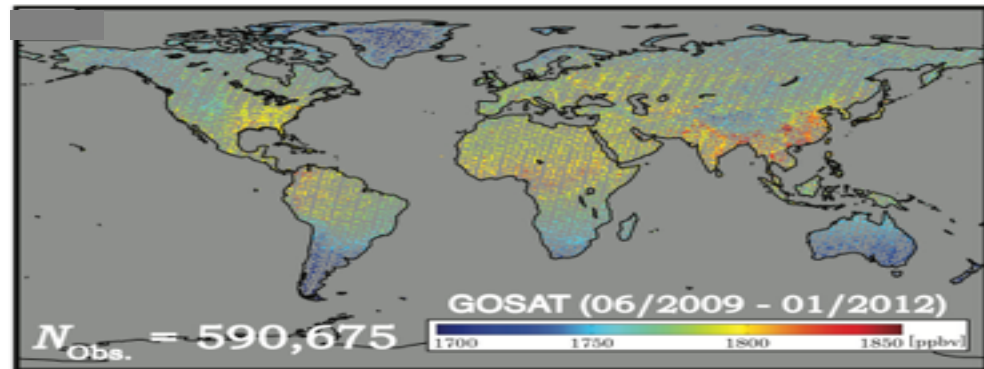
CO, methane, other volatile organic compounds (VOCs)

- **CO observed from space: 50-200 ppb**

Sources: fuel combustion, open fires, VOC oxidation

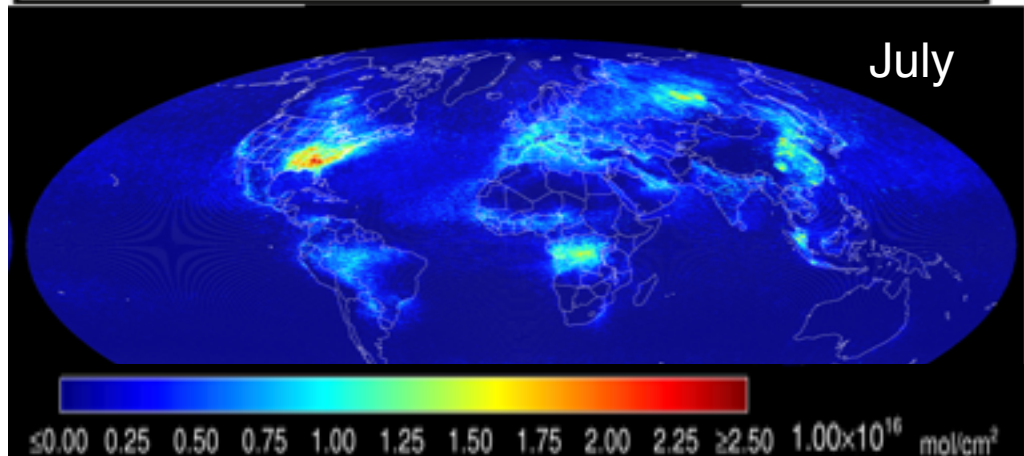


- **Methane observed from space: 1700-1900 ppb**
Sources: wetlands, livestock, oil/gas production, landfills, coal mines...



- **Formaldehyde (HCHO) observed from space: product of VOC oxidation**

Non-methane VOC sources: vegetation, fuel combustion, open fires, industry



Methylchloroform as a proxy for global tropospheric OH

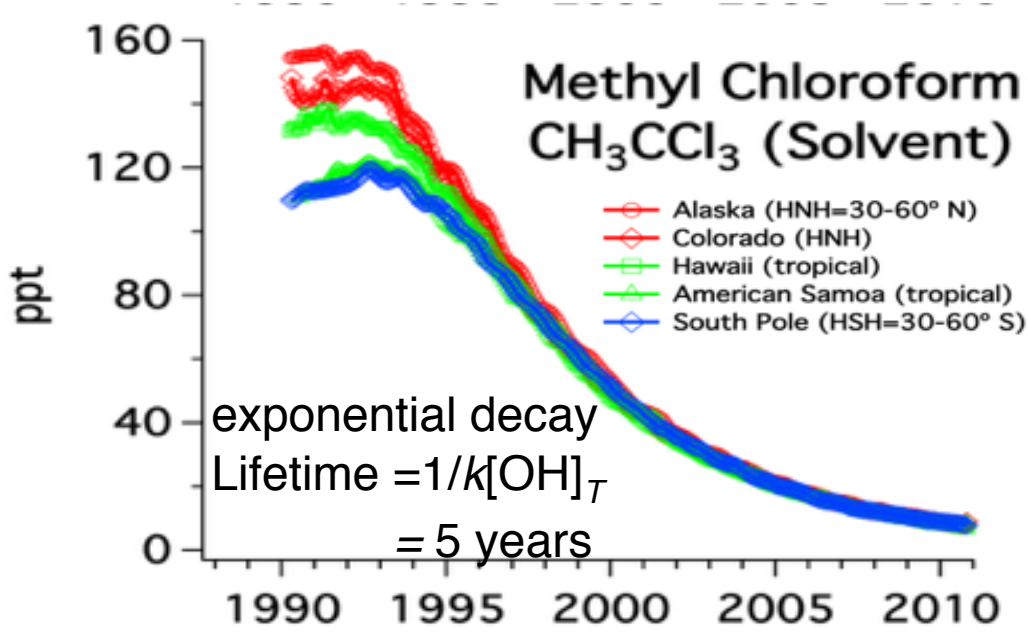
Methylchloroform (CH_3CCl_3):

- Uniquely anthropogenic (industrial solvent), banned by Montreal Protocol
- Removed from troposphere by oxidation by OH, transport to troposphere

Mass balance equation for troposphere (T):

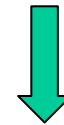
Observed decay rate $\frac{dm_T}{dt} = -km_T [\text{OH}]_T - k_{TS}m_T + k_{ST}m_S$

loss rate from $\text{CH}_3\text{CCl}_3 + \text{OH}$
exchange with stratosphere (small)



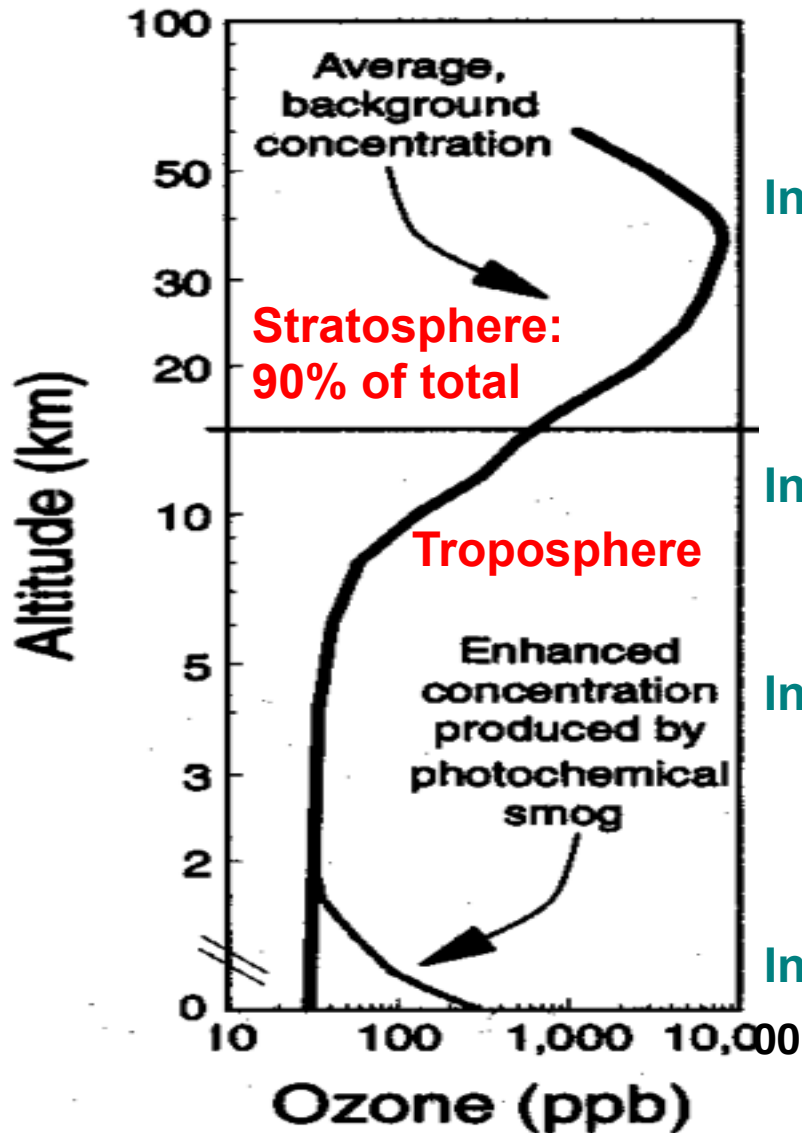
$$[\text{OH}]_T = \frac{-dm_T / dt}{km_T}$$

$$= 1.0 \times 10^6 \text{ molecules cm}^{-3}$$



$\tau_{\text{CO}} = 2 \text{ months}$
 $\tau_{\text{CH}_4} = 10 \text{ years}$

Several reasons to care about tropospheric ozone



In stratosphere: UV shield

Stratosphere:
90% of total

In middle/upper troposphere: greenhouse gas

Troposphere

In troposphere: source of OH

Enhanced
concentration
produced by
photochemical
smog

In surface air: toxic to humans and vegetation

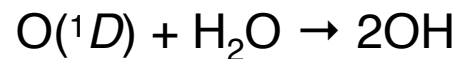
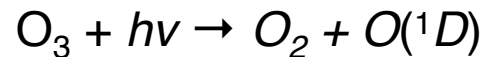
Closing the tropospheric ozone budget requires a large chemical source in troposphere

Transport from stratosphere



500 Tg O₃ yr⁻¹

Tropopause



—————▶ Chemical loss

2200 Tg O₃ yr⁻¹

Deposition



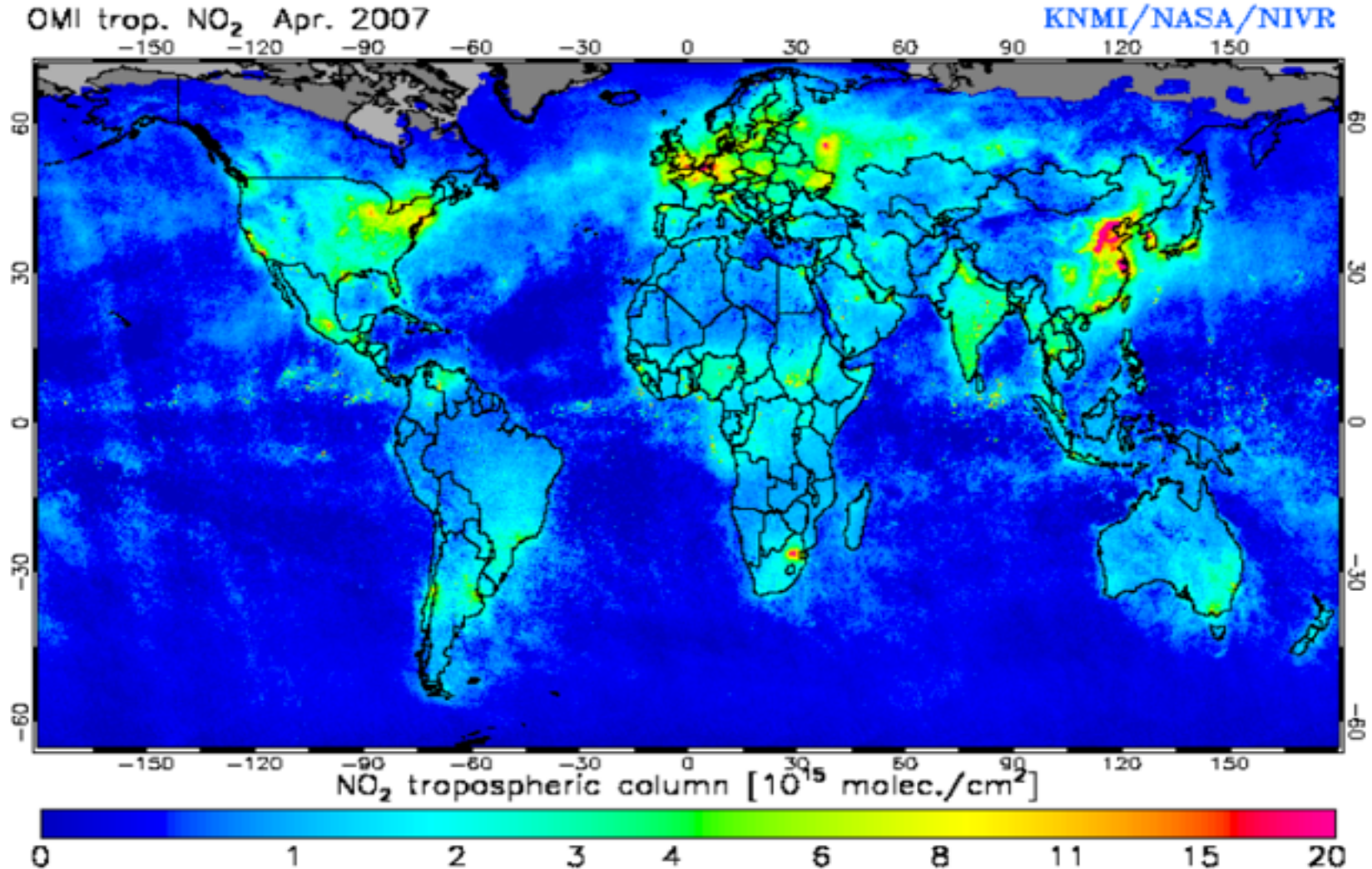
1000 Tg O₃ yr⁻¹

A large tropospheric source is needed to balance the loss

Tropospheric NO_x: critical agent for ozone production

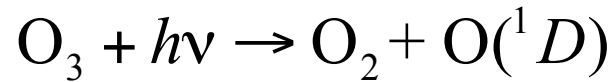
Sources: fuel combustion, open fires, lightning, soils

Tropospheric NO₂ columns measured from satellite

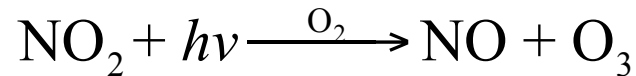
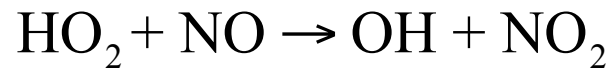
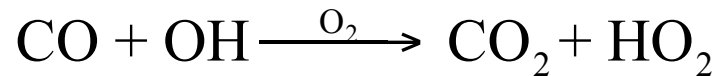


Production of tropospheric ozone by oxidation of CO and VOCs catalyzed by HO_x and NO_x radicals

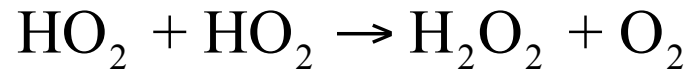
Initiation: production of HO_x radicals



Propagation: oxidation of CO (simplest case) in presence of NO_x



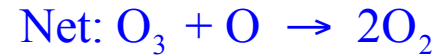
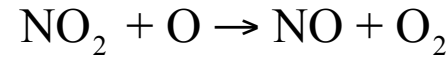
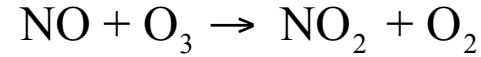
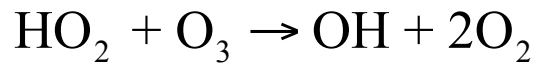
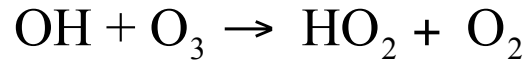
Termination: loss of HO_x radicals



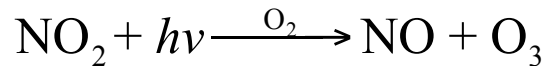
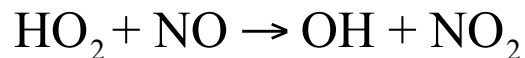
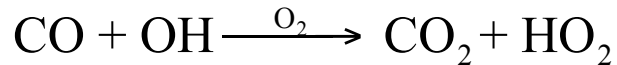
followed by H₂O₂ and HNO₃ deposition

HO_x and NO_x radicals catalyze ozone production in troposphere, but loss in the stratosphere – why the difference?

In stratosphere,

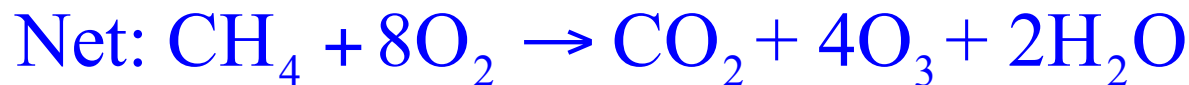
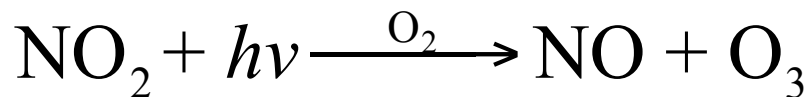
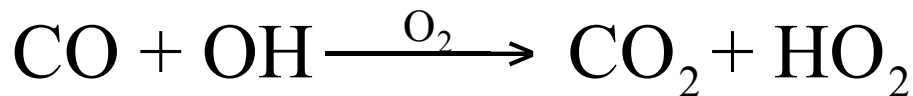
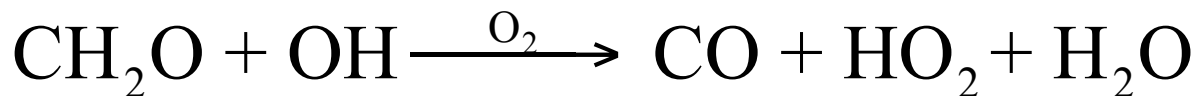
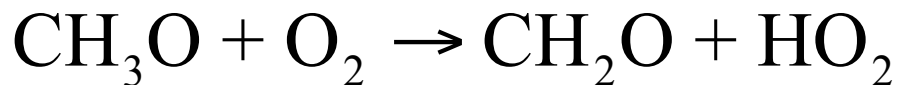
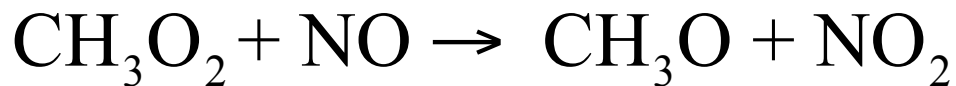
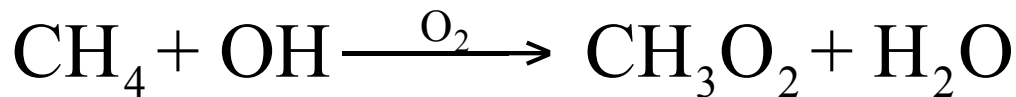


In troposphere,

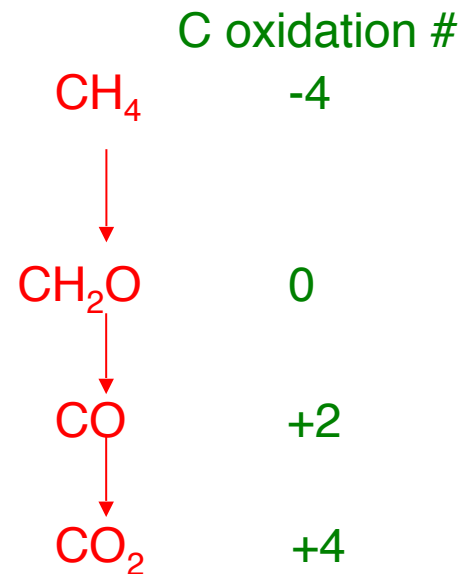


- CO/O₃ and NO/O₃ ratios are much higher in troposphere than stratosphere
- O concentrations are very low in troposphere

Methane oxidation cascade follows same schematic as CO



- Expanded HO_x family: HO_x ≡ OH + HO₂ + CH₃O₂ + CH₃O
- Oxidation of methane by above mechanism produces four ozone molecules



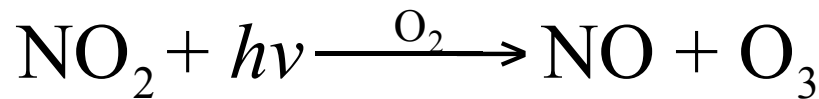
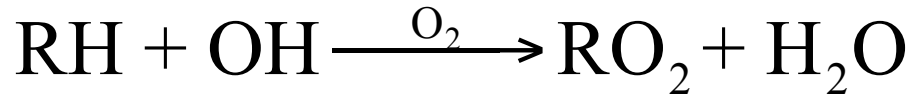
Questions

1. How do you expect recovery of the stratospheric ozone layer to affect tropospheric OH concentrations?
2. Maximum photon flux during summer results in a seasonal maximum of ozone in polluted regions but a seasonal minimum of ozone in very clean regions. Why is that?

Same basic mechanism for higher VOCs but many branches

RO₂ can also isomerize, or react with HO₂ or RO₂, producing peroxides, epoxides, alcohols, carboxylic acids...

OH can also add to double bonds of unsaturated VOCs, producing hydroxyorganics



RO can also decompose or isomerize to produce a range of aldehydes, ketones, dicarbonyls...

Oxidation product goes on to react with OH, adding functionality and making more ozone

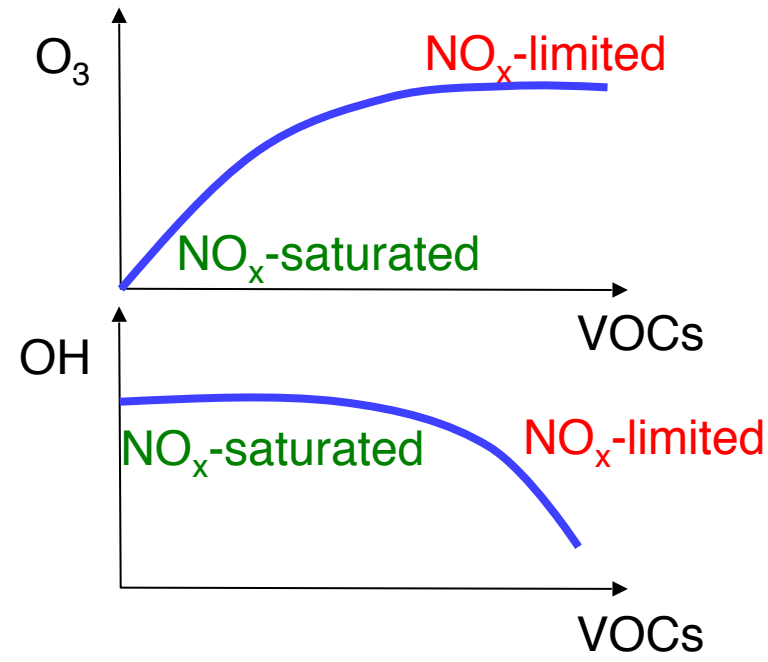
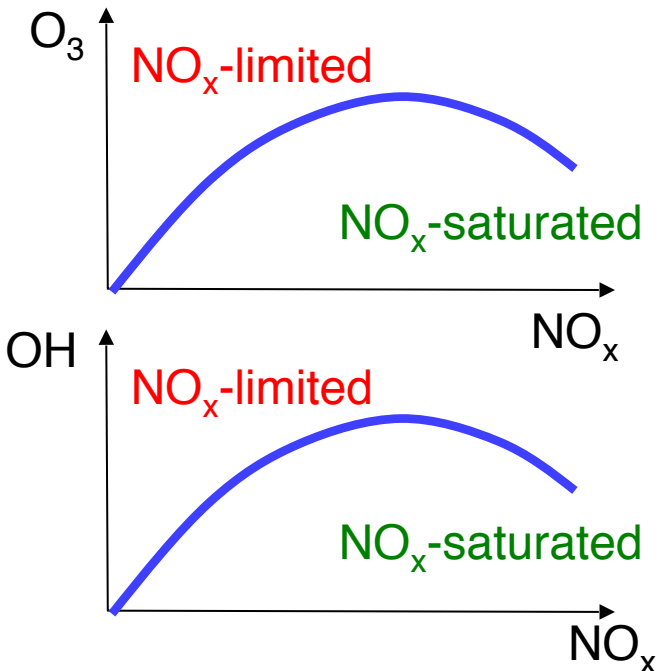
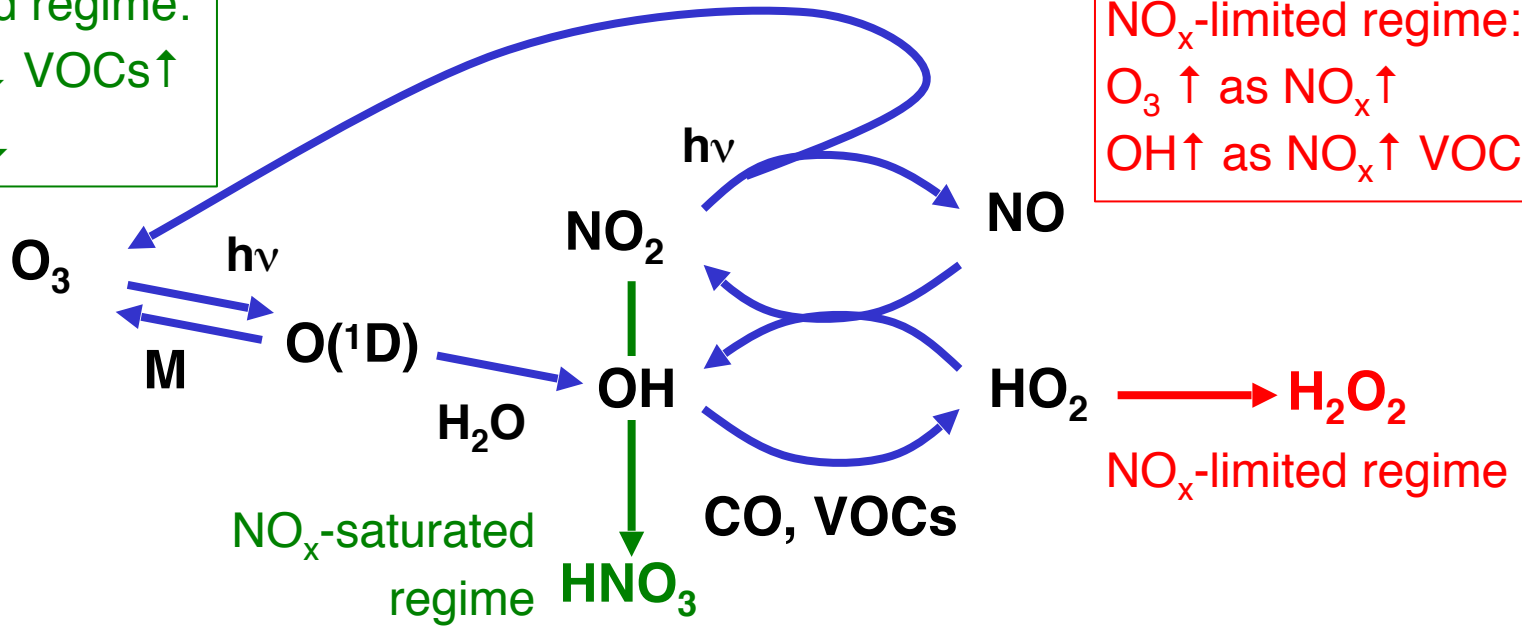
A brief cheat sheet for organic functions

RCHO	aldehyde
RC(O)R'	ketone
RC(O)-	acyl group
RC(O)C(O)R'	dicarbonyl
RO ₂	peroxy radical
ROOR'	peroxide
ROOH	hydroperoxide
ROH	hydroxy, alcohol
RC(O)CR'	epoxide (cyclic ether)
RC(O)OH	carboxylic acid
RC(O)OR'	ester
RONO ₂	nitrate

Dependence of ozone production and OH on NO_x and VOCs

NO_x -saturated regime:
 $\text{O}_3 \uparrow$ as $\text{NO}_x \downarrow$ VOCs \uparrow
 $\text{OH} \uparrow$ as $\text{NO}_x \downarrow$

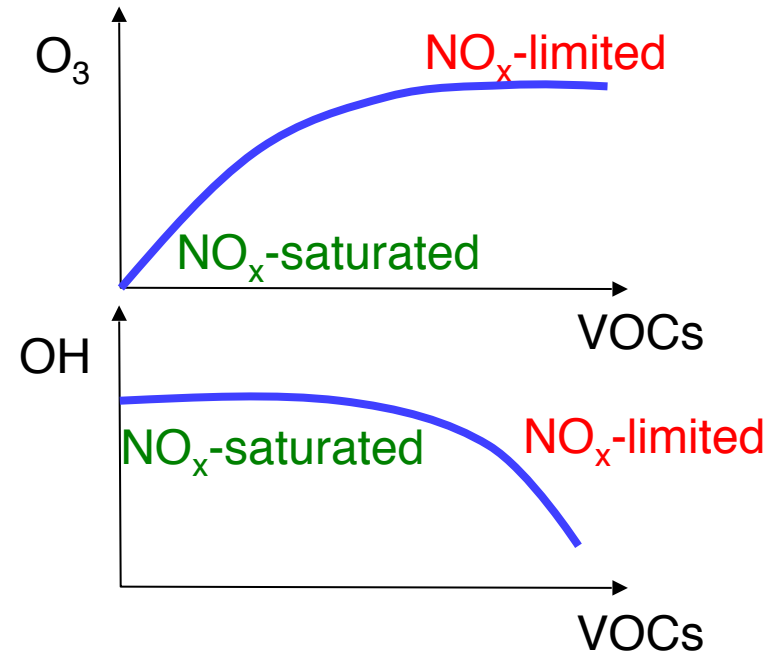
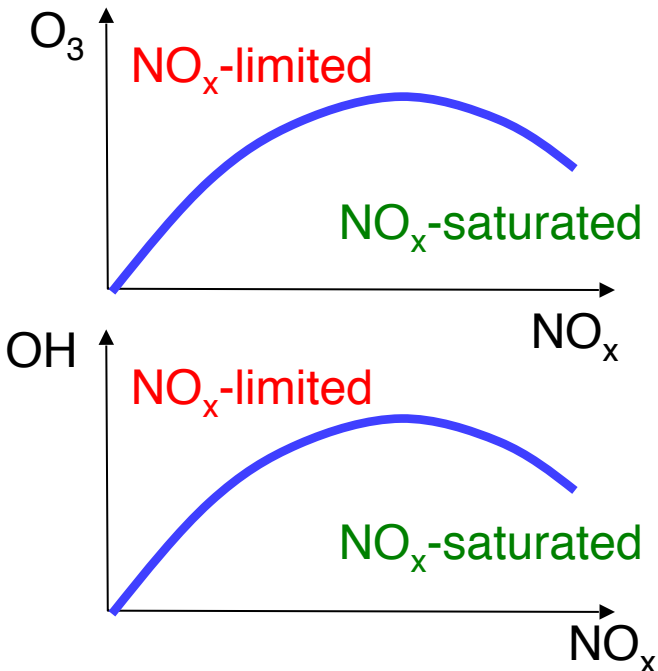
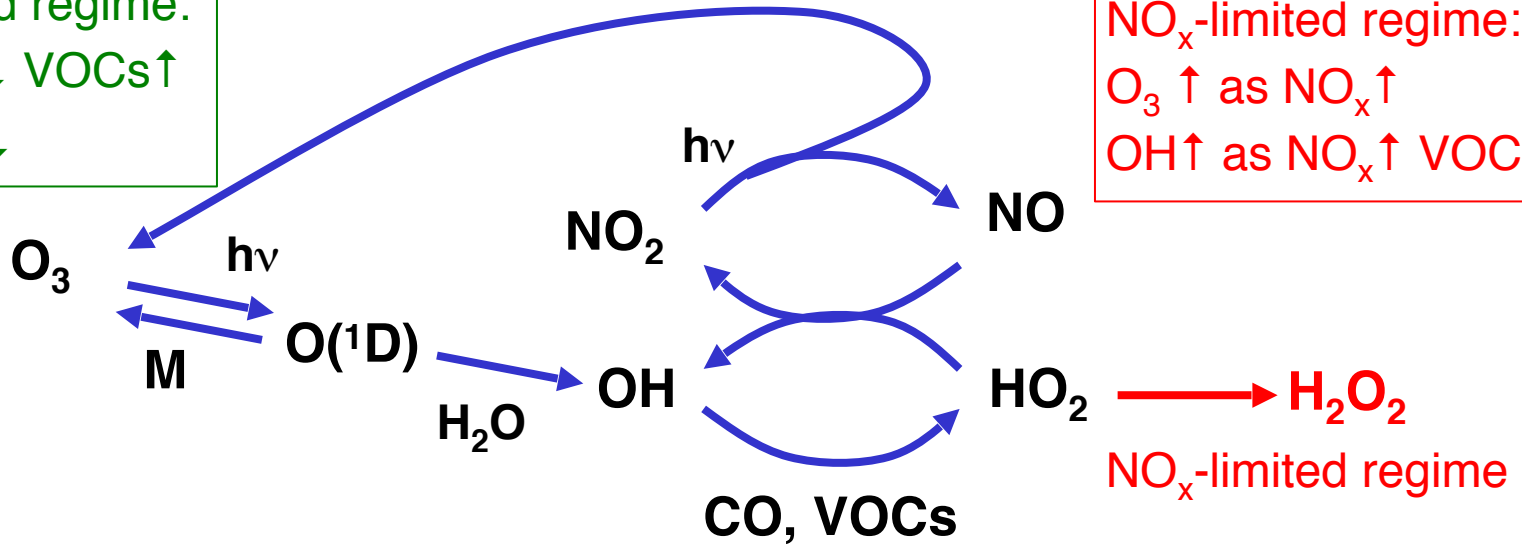
NO_x -limited regime:
 $\text{O}_3 \uparrow$ as $\text{NO}_x \uparrow$
 $\text{OH} \uparrow$ as $\text{NO}_x \uparrow$ VOCs \downarrow



Dependence of ozone production and OH on NO_x and VOCs

NO_x-saturated regime:
O₃ ↑ as NO_x ↓ VOCs ↑
OH ↑ as NO_x ↓

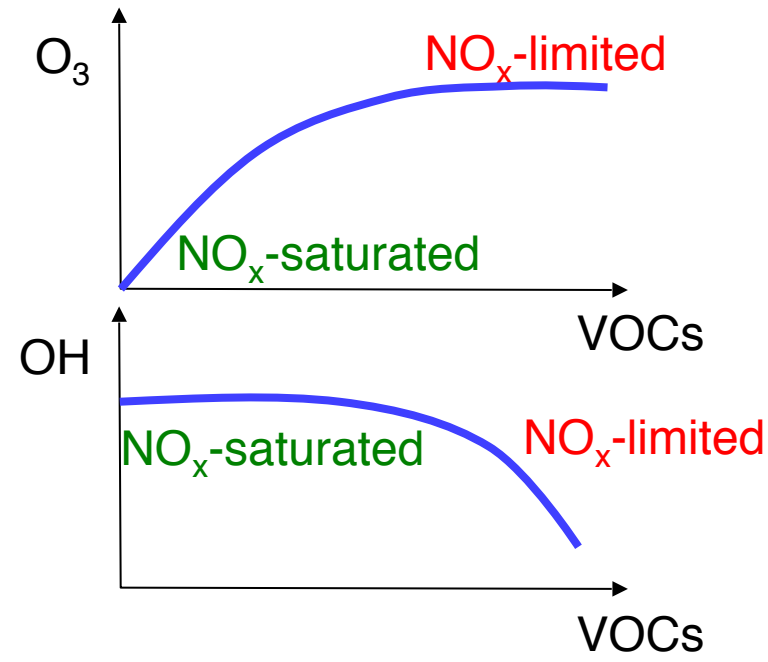
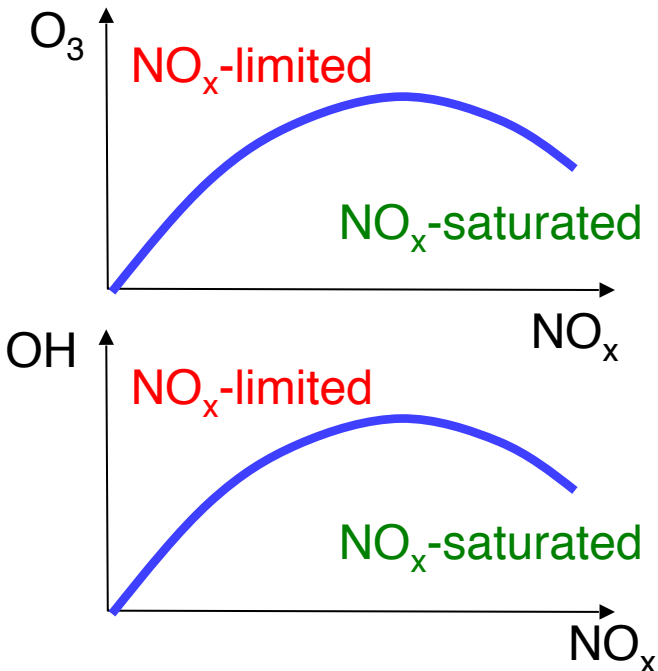
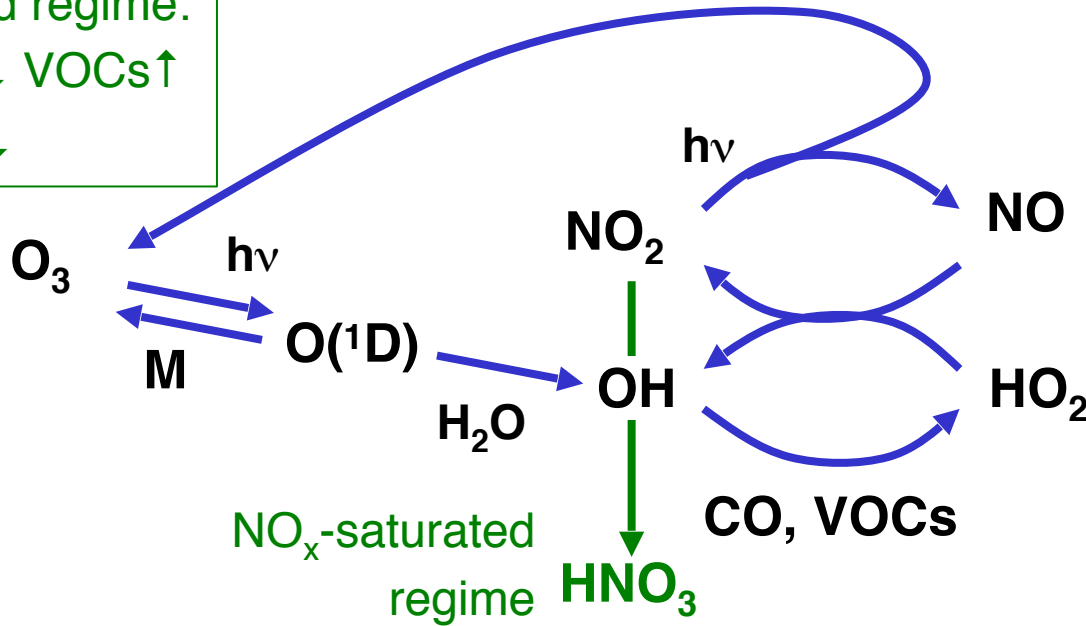
NO_x-limited regime:
O₃ ↑ as NO_x ↑
OH ↑ as NO_x ↑ VOCs ↓



Dependence of ozone production and OH on NO_x and VOCs

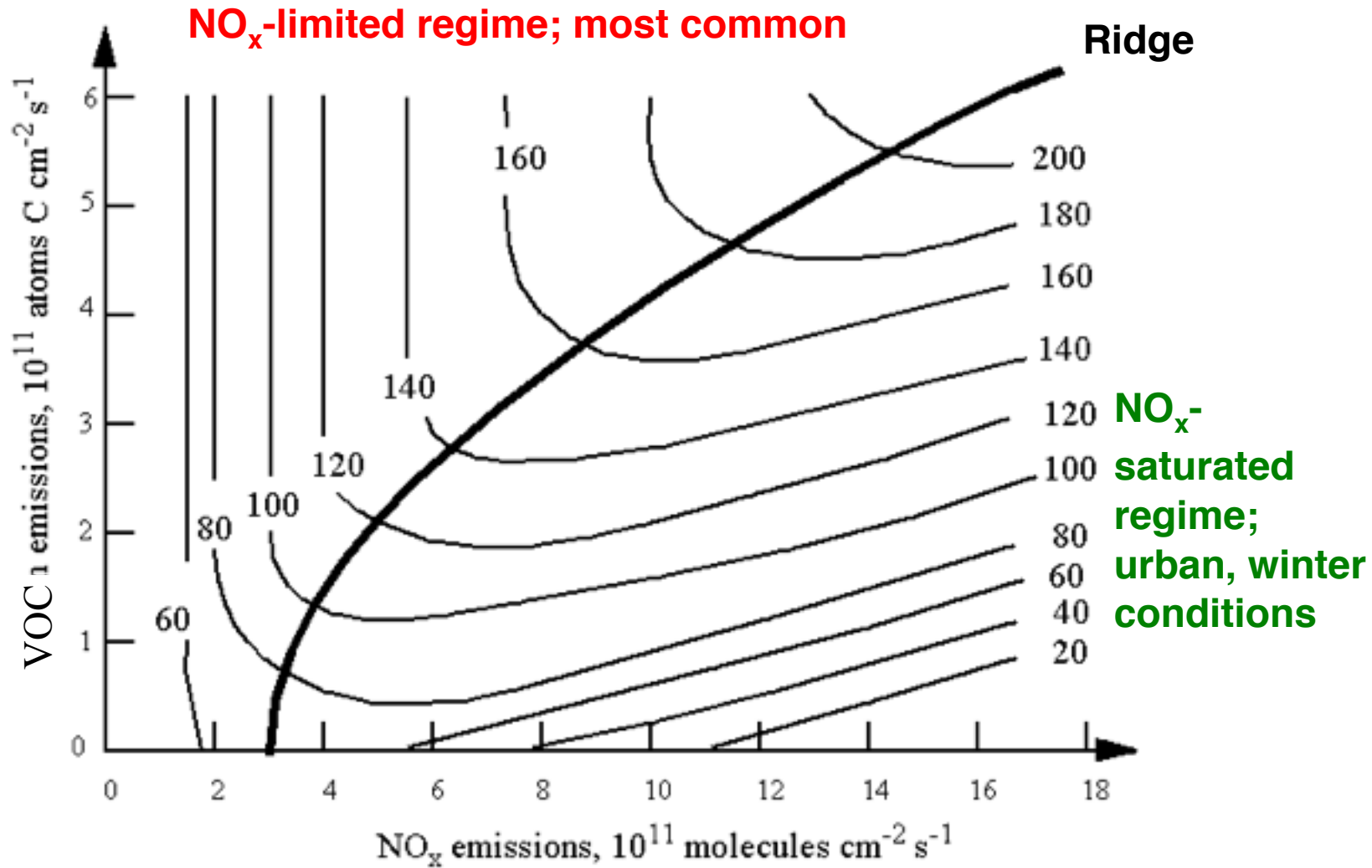
NO_x -saturated regime:
 $\text{O}_3 \uparrow$ as $\text{NO}_x \downarrow$ VOCs \uparrow
 $\text{OH} \uparrow$ as $\text{NO}_x \downarrow$

NO_x -limited regime:
 $\text{O}_3 \uparrow$ as $\text{NO}_x \uparrow$
 $\text{OH} \uparrow$ as $\text{NO}_x \uparrow$ VOCs \downarrow



OZONE CONCENTRATIONS vs. NO_x AND VOC EMISSIONS

Sample box model calculation with detailed mechanism

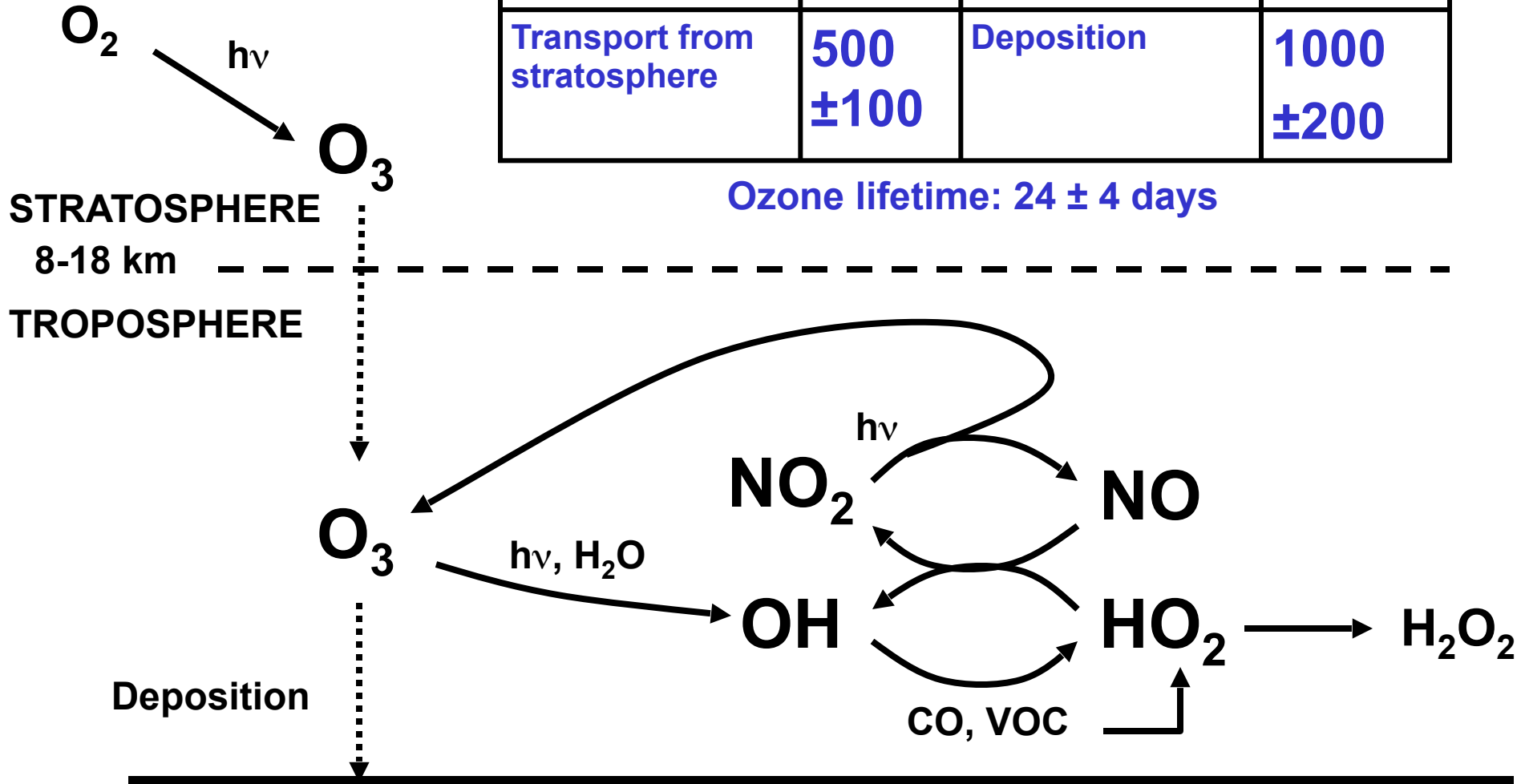


GLOBAL BUDGET OF TROPOSPHERIC OZONE ($\text{Tg O}_3 \text{ yr}^{-1}$)

IPCC (2007) average of 12 models

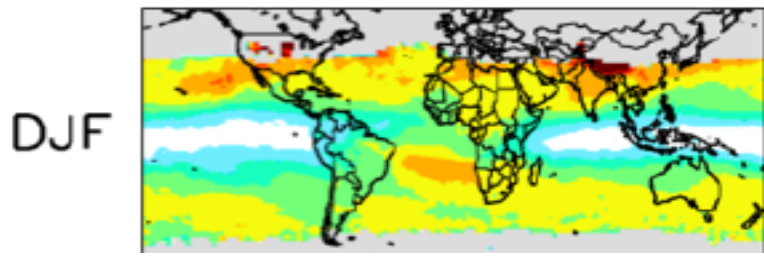
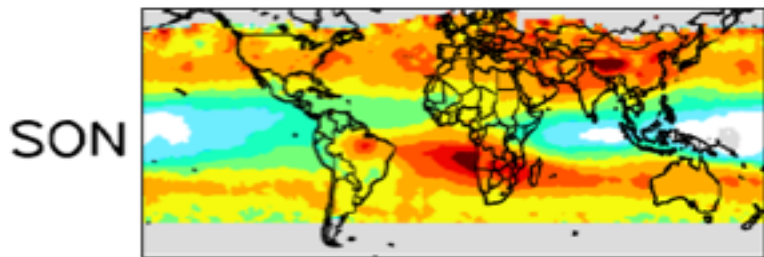
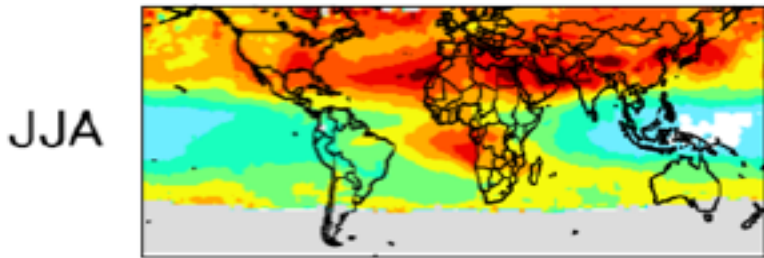
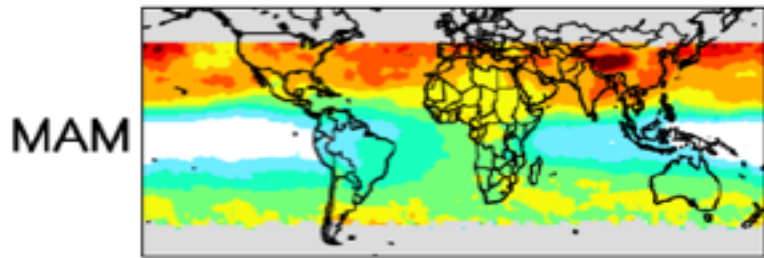
Chem prod in troposphere	4700 ± 700	Chem loss in troposphere	4200 ± 500
Transport from stratosphere	500 ± 100	Deposition	1000 ± 200

Ozone lifetime: 24 ± 4 days



GLOBAL DISTRIBUTION OF TROPOSPHERIC OZONE

OMI satellite observations at 700-400 hPa



25 30 35 40 45 50 55 60 65 70 Ozone, ppbv

- Maximum values at northern mid-latitudes in spring-summer due to anthropogenic pollution;

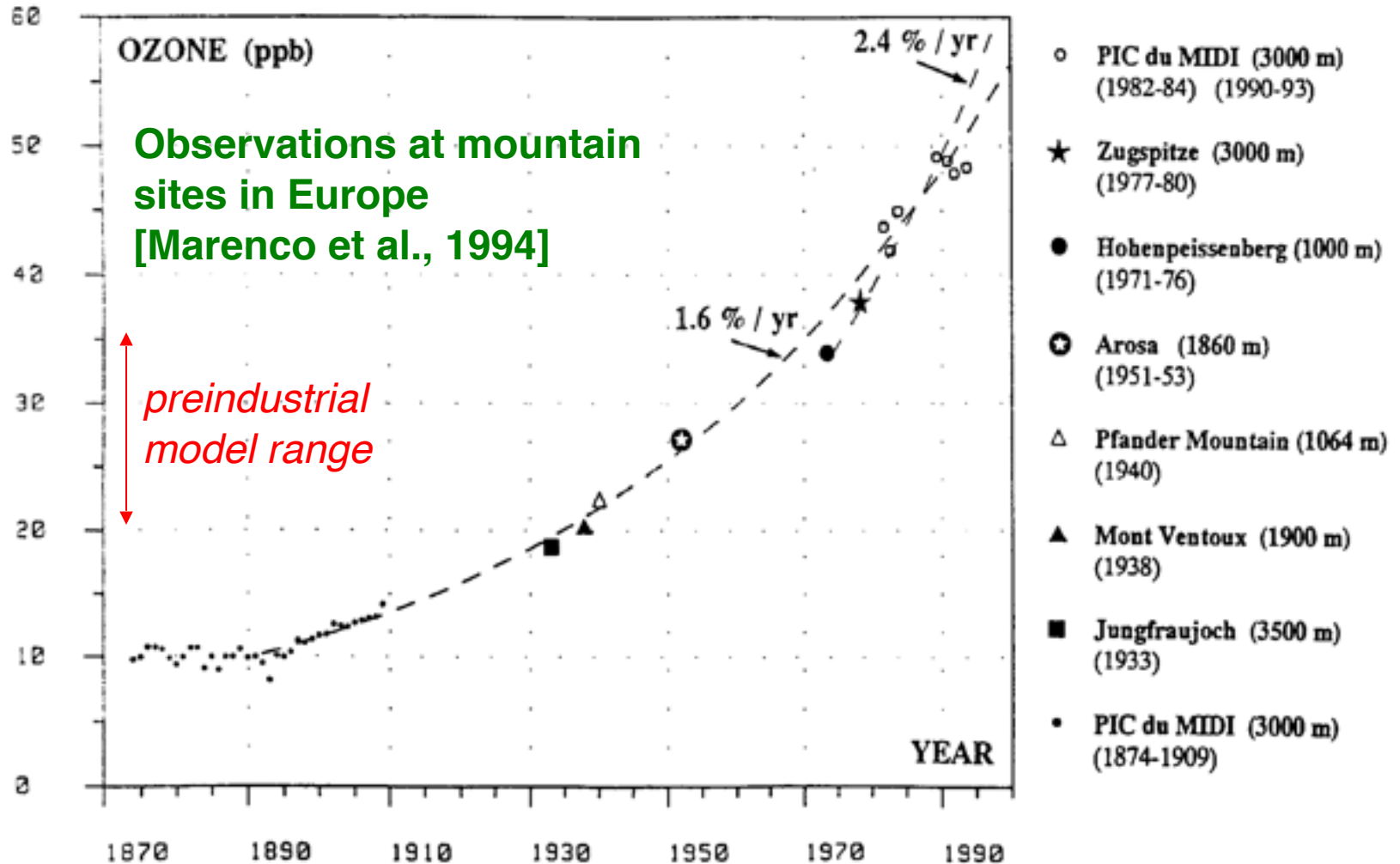
- High values in tropical regions affected by seasonal biomass burning;

- Minimum values over tropical oceans due to chemical loss

Hu et al. [2017]

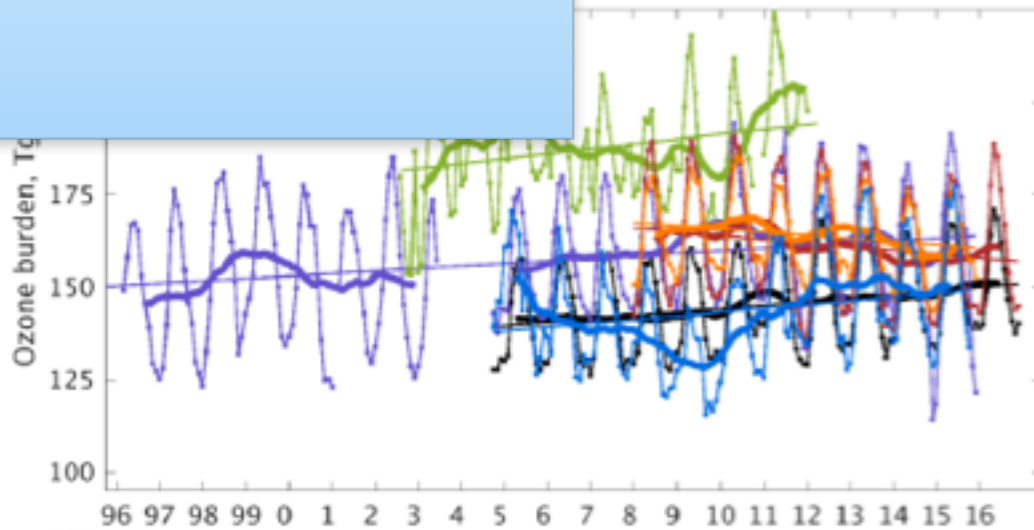
Observed rise of tropospheric ozone in 20th century

thought to be due mostly to anthropogenic NO_x and methane but this is uncertain

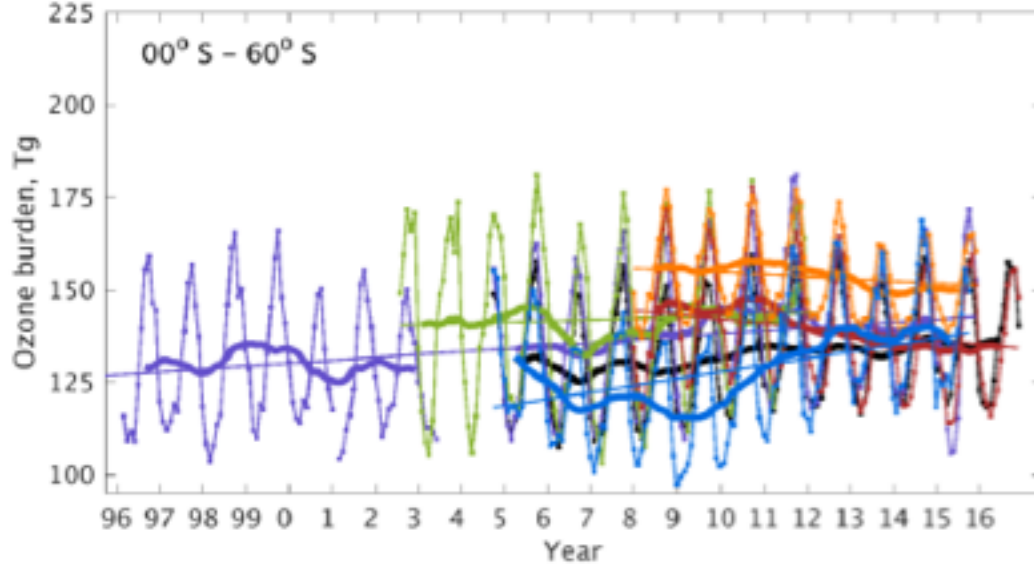


tropospheric ozone seen from satellites

rise in ozone that models cannot explain

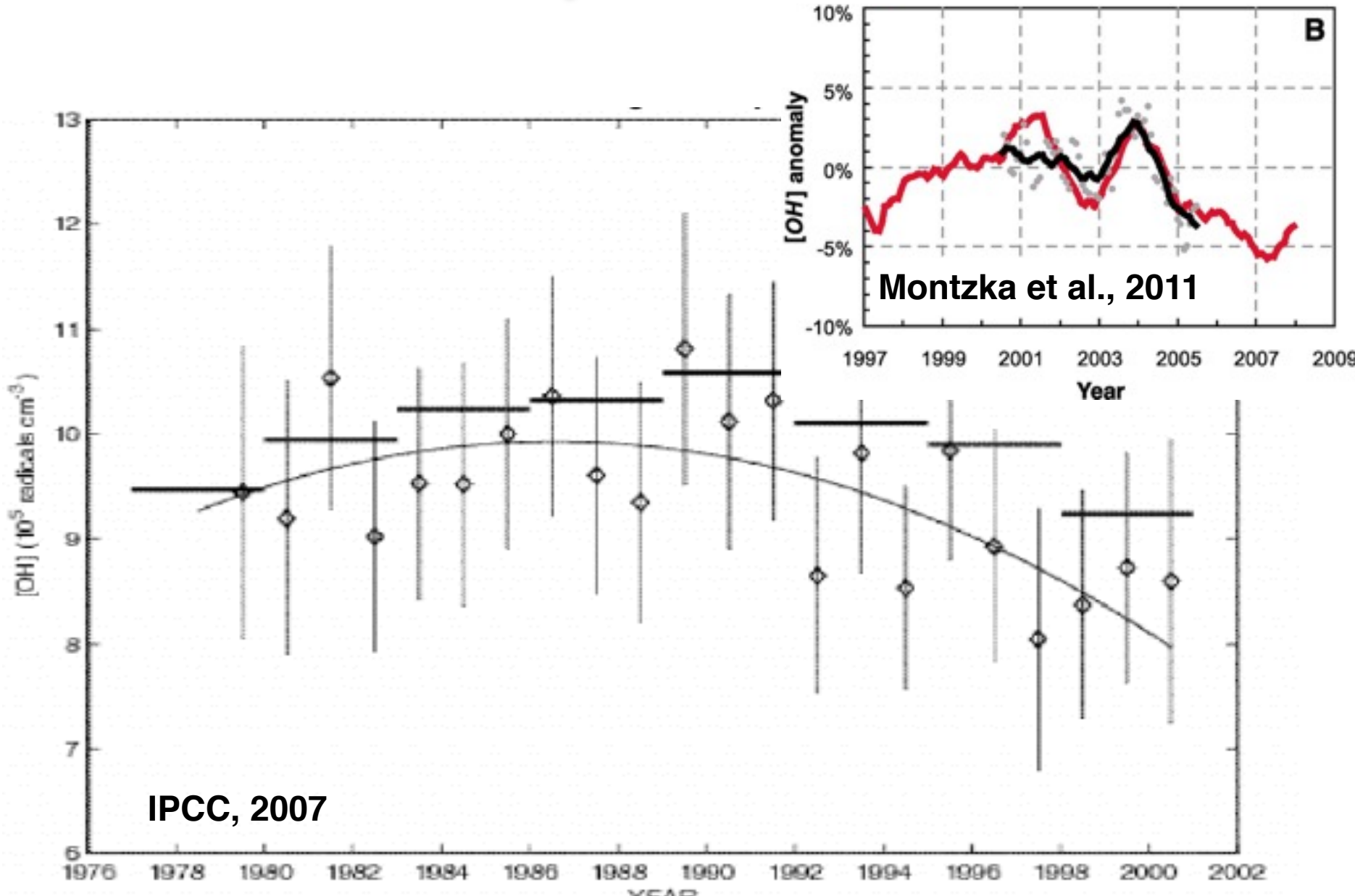


	change, Tg yr ⁻¹	p-value
Black: OMI/MLS	0.95 +/- 0.55	0.00
Brown: IASI-FORLI	-1.01 +/- 1.17	0.09
Orange: IASI-SOFRID	-0.86 +/- 0.95	0.07
Purple: GOME/OMI	0.68 +/- 0.39	0.00
Blue: OMI-RAL	1.02 +/- 0.79	0.01
Green: SCIAMACHY	1.33 +/- 1.04	0.01



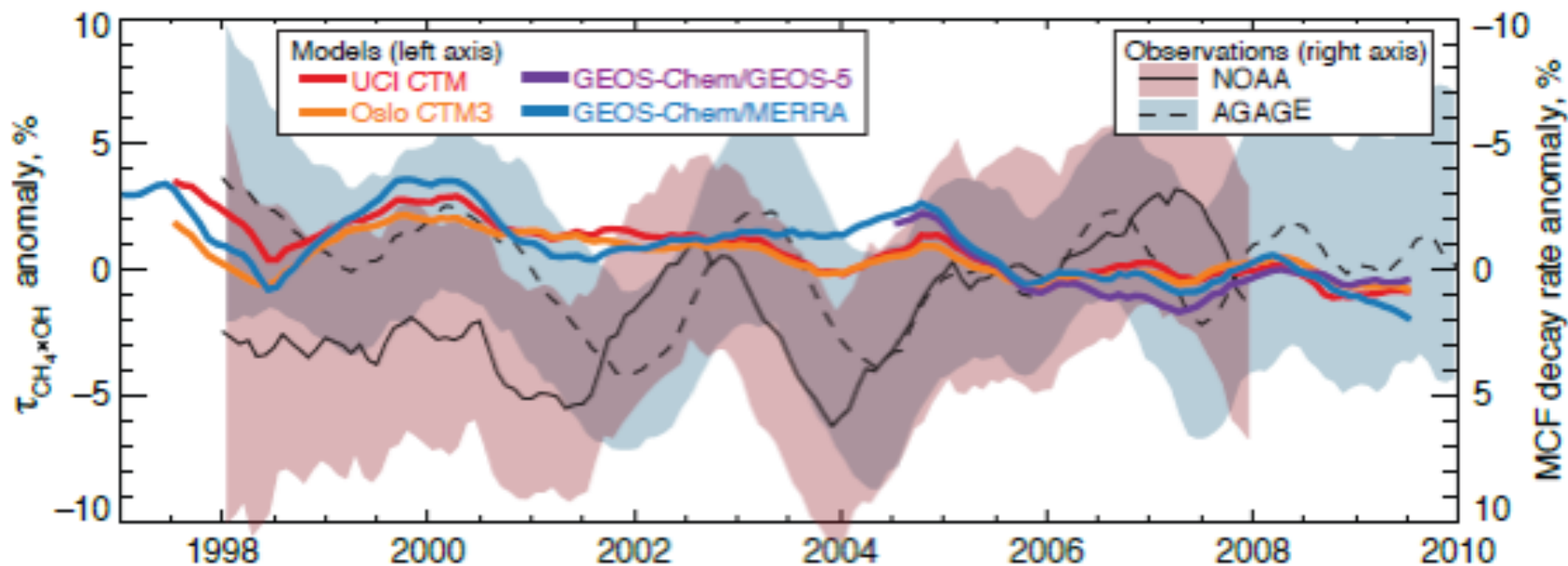
	change, Tg yr ⁻¹	p-value
Black: OMI/MLS	0.83 +/- 0.64	0.01
Brown: IASI-FORLI	-1.14 +/- 1.14	0.05
Orange: IASI-SOFRID	-0.56 +/- 0.97	0.26
Purple: GOME/OMI	0.80 +/- 0.40	0.00
Blue: OMI-RAL	1.83 +/- 0.84	0.00
Green: SCIAMACHY	0.34 +/- 1.48	0.65

TRENDS IN GLOBAL TROPOSPHERIC OH inferred from methylchloroform observations



But current models have no skill in simulating observed OH trends

Lightning NO_x is the dominant factor of interannual variability in models



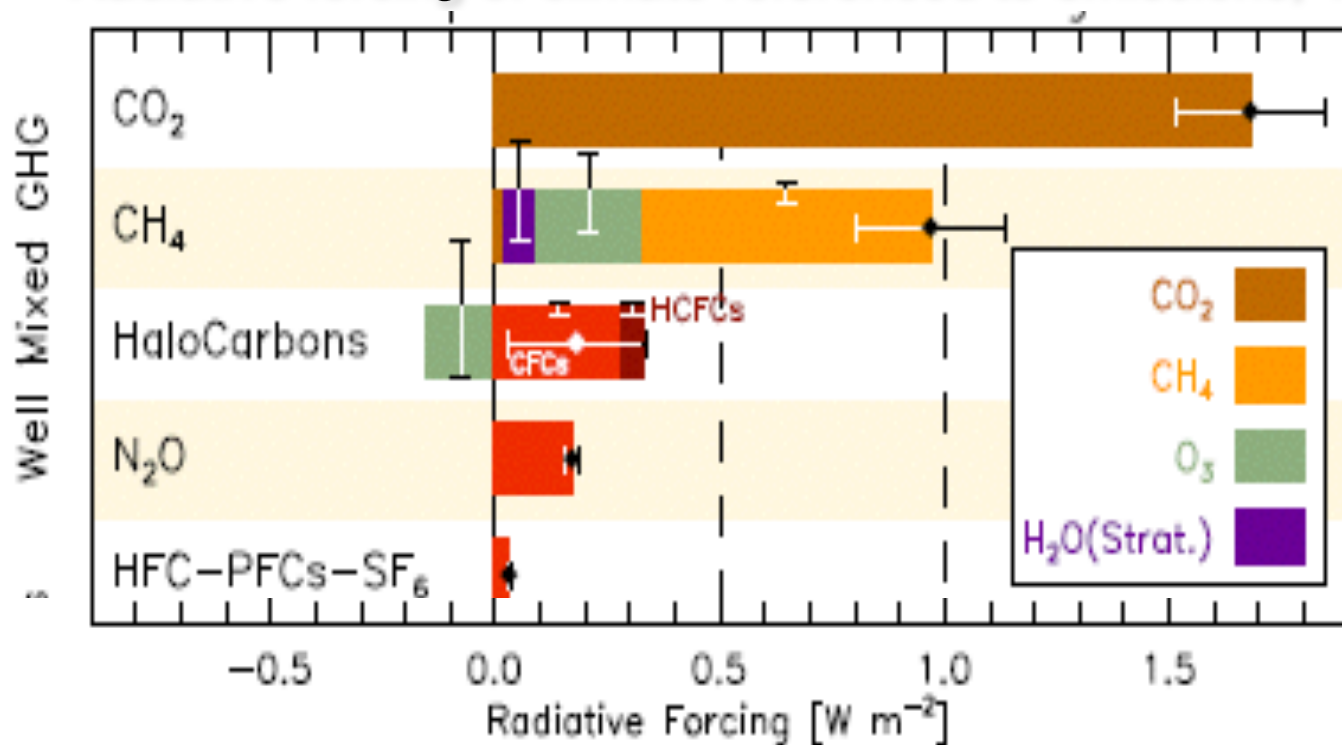
Better understanding of OH changes is needed!

Questions

1. If the methane source to the atmosphere were to double, would the methane concentration (a) double, (b) less than double, (c) more than double?
2. Atmospheric measurements of the $\text{H}_2\text{O}_2/\text{HNO}_3$ concentration ratio offer a simple diagnostic of whether ozone production in a polluted environment is NO_x -limited or NO_x -saturated. Why?

Methane: 2nd anthropogenic greenhouse gas after CO₂

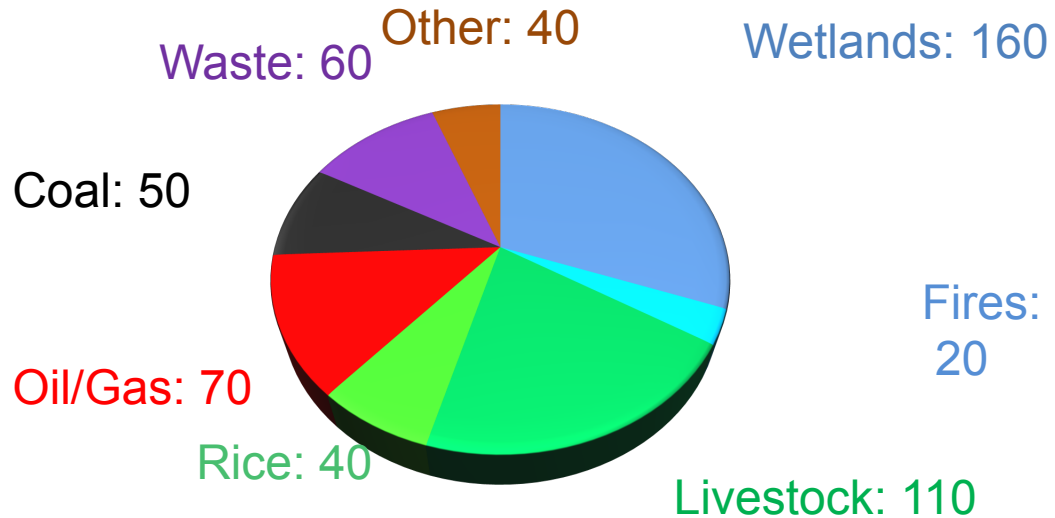
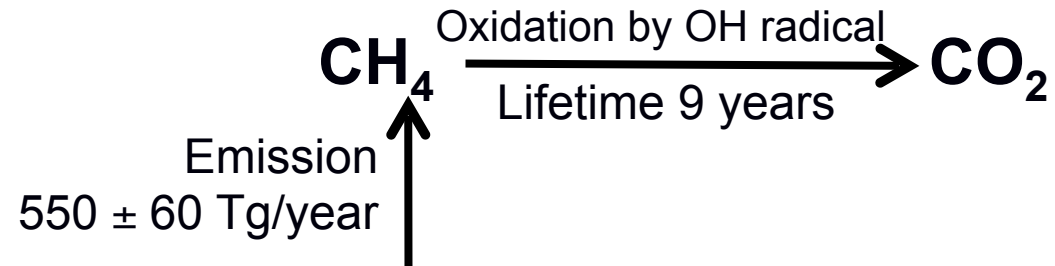
Radiative forcing of climate referenced to emissions, 1750-2011



[IPCC, 2014]

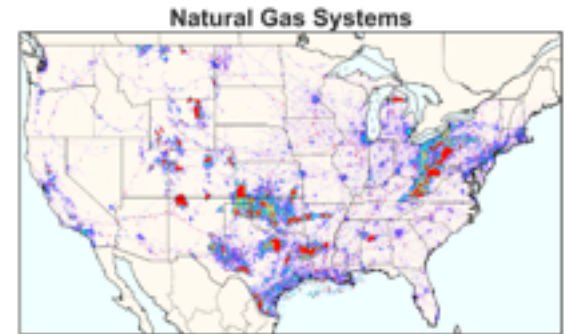
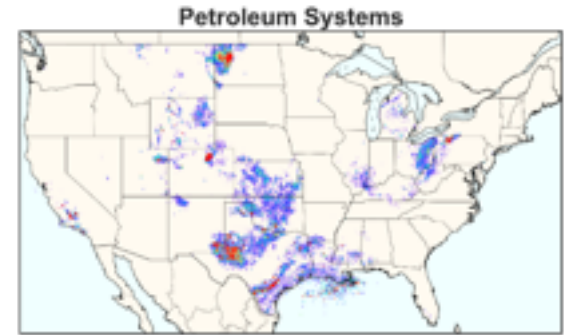
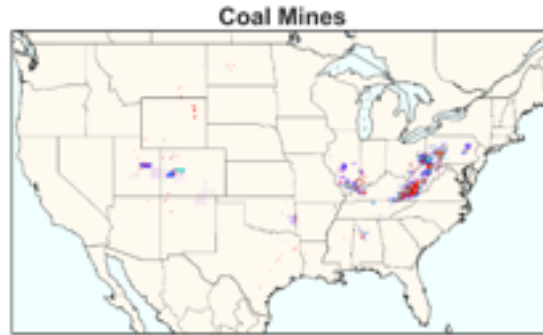
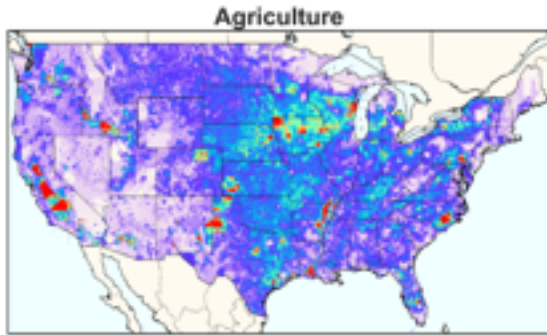
- Methane is 60% as important as CO₂ in explaining past warming
- Atmospheric lifetime of methane is 9 years, much shorter than CO₂ (> 100 years)
- Methane is most relevant as a near-term (~20 years) climate forcer
- Methane and CO₂ emissions should not be “equivalent” in climate policy
- Reducing methane emissions has air quality co-benefits and can make money

Complexity of methane sources

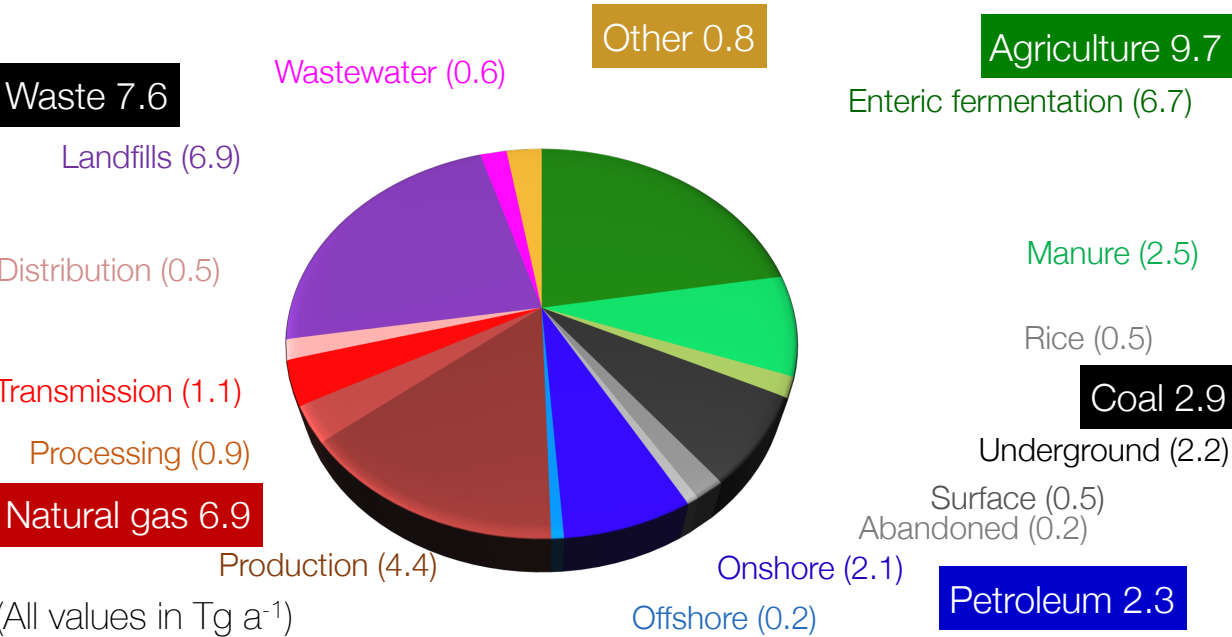


Global totals (Tg/y), EDGAR4.2 and LPJ wetlands

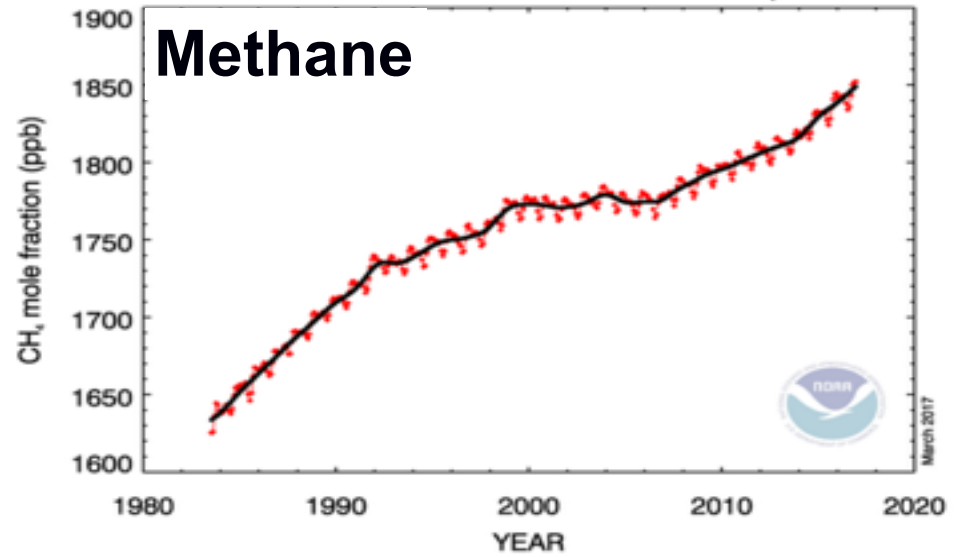
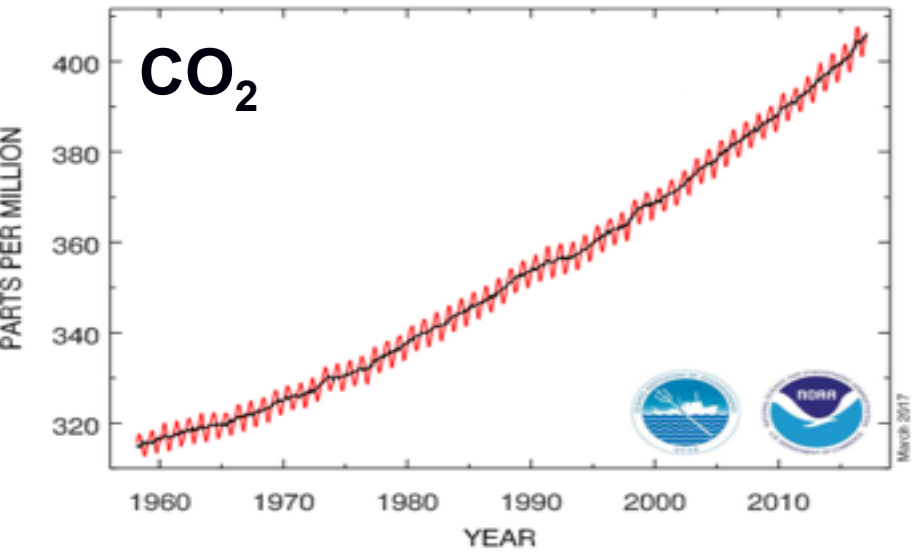
EPA anthropogenic methane inventory for the US



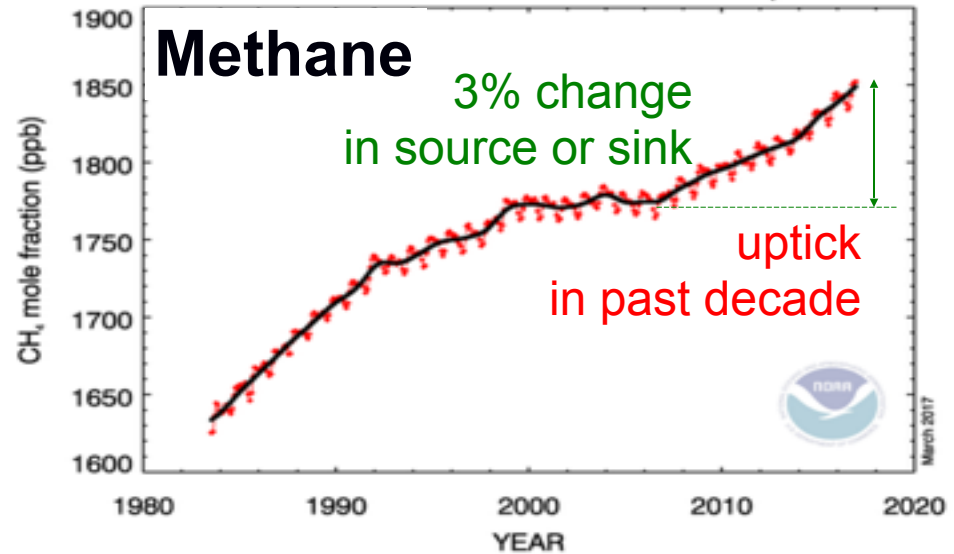
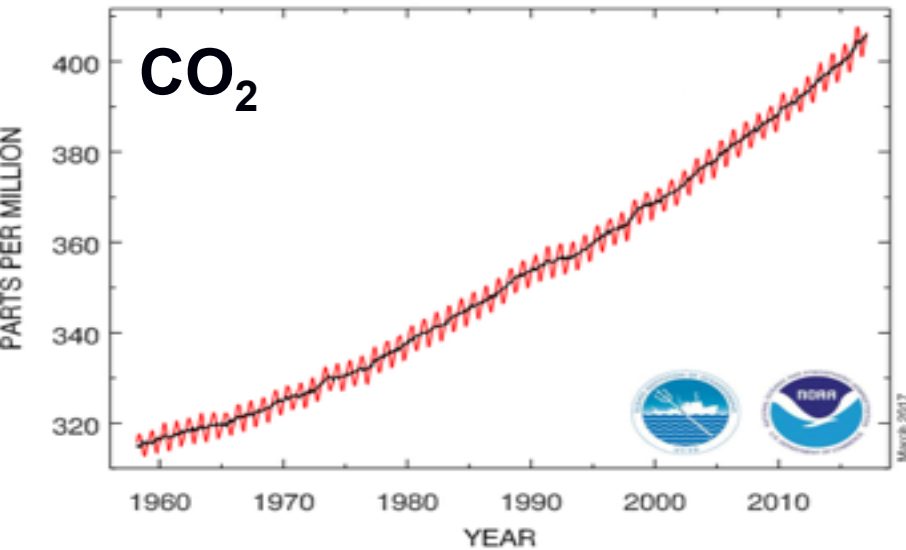
EPA US inventory (2012) submitted to UNFCCC



Methane fits and starts over past 40 years



Methane fits and starts over past 40 years



- Uptick since 2006 has been attributed to oil/gas, livestock, wetlands, OH...
- ...but it represents only 3% change in sources or sinks, on top of a factor 2.5 increase since preindustrial times that we don't understand!

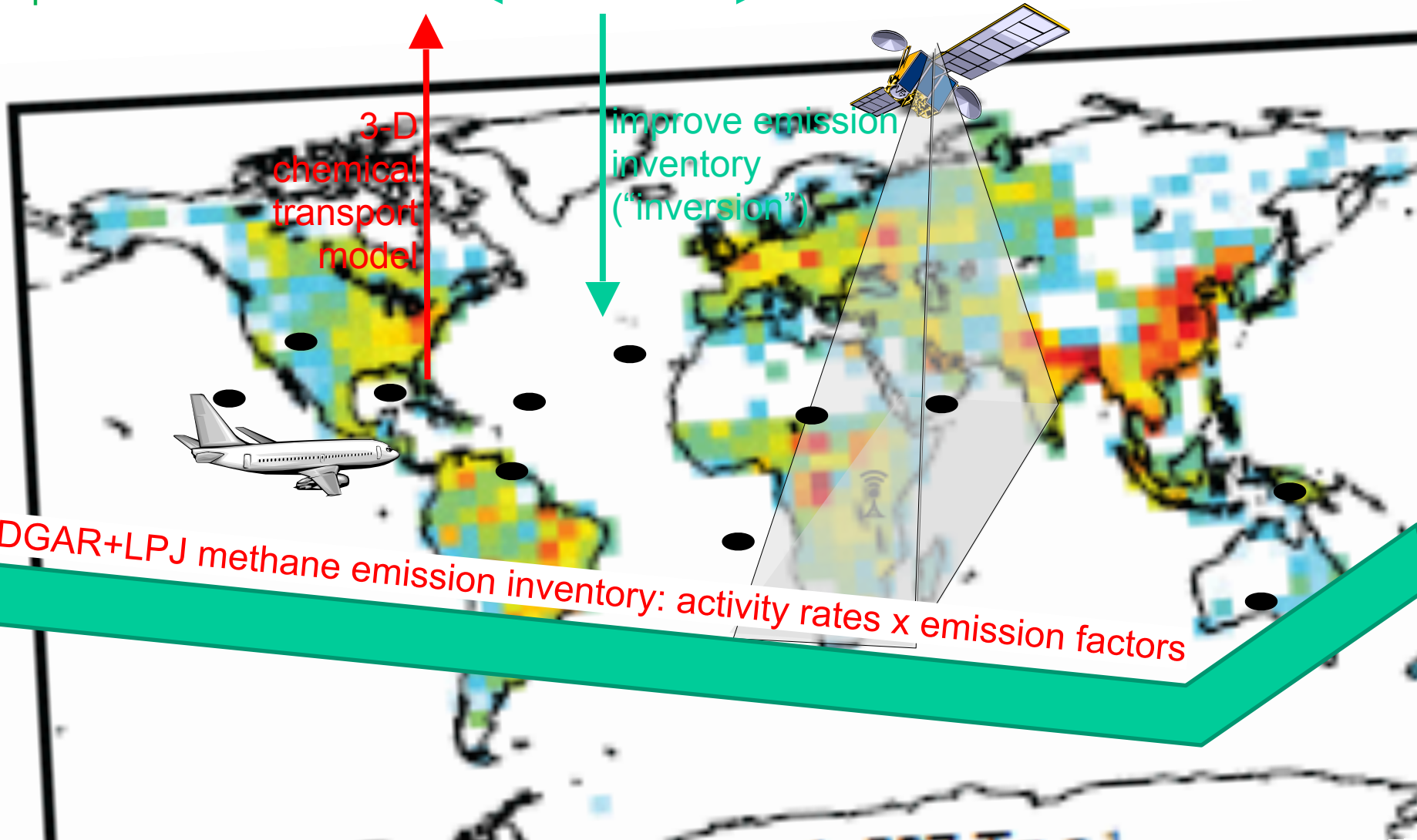
Using atmospheric methane observations to improve emission inventories

predicted concentrations ← compare → observed concentrations

3-D
chemical
transport
model

improve emission
inventory
("inversion")

EDGAR+LPJ methane emission inventory: activity rates x emission factors



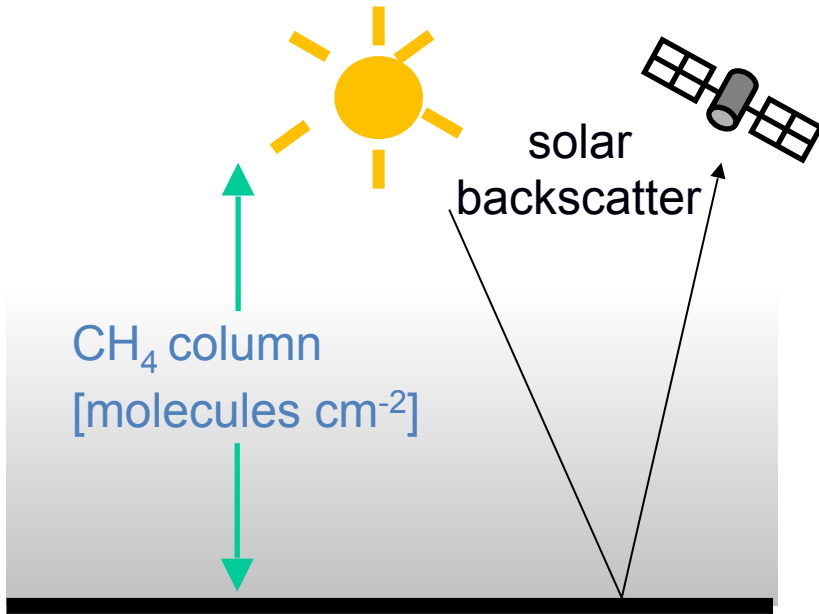
Sometimes methane is easy to observe by remote sensing...

Aliso Canyon, CA methane leak

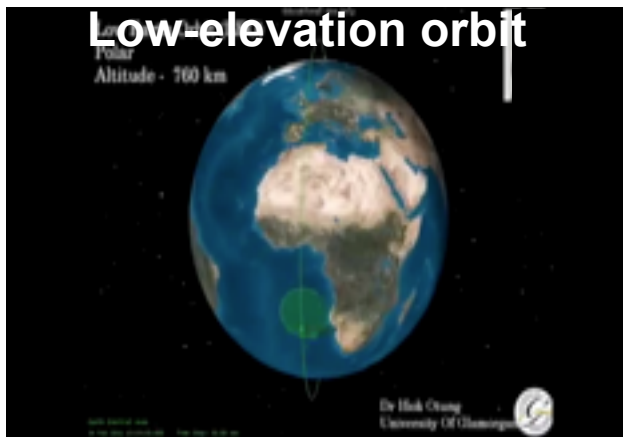


...but generally it's much harder!

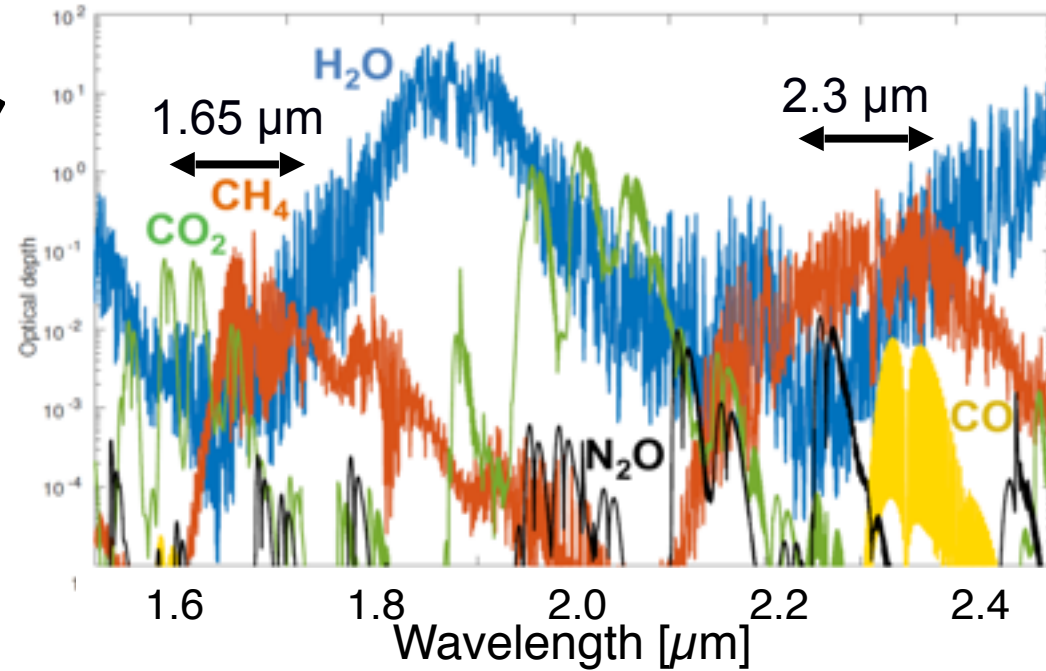
Observing methane atmospheric columns from space



GOSAT (JAXA), 2009- ;
TROPOMI (ESA), 2017 launch



Atmospheric optical depths

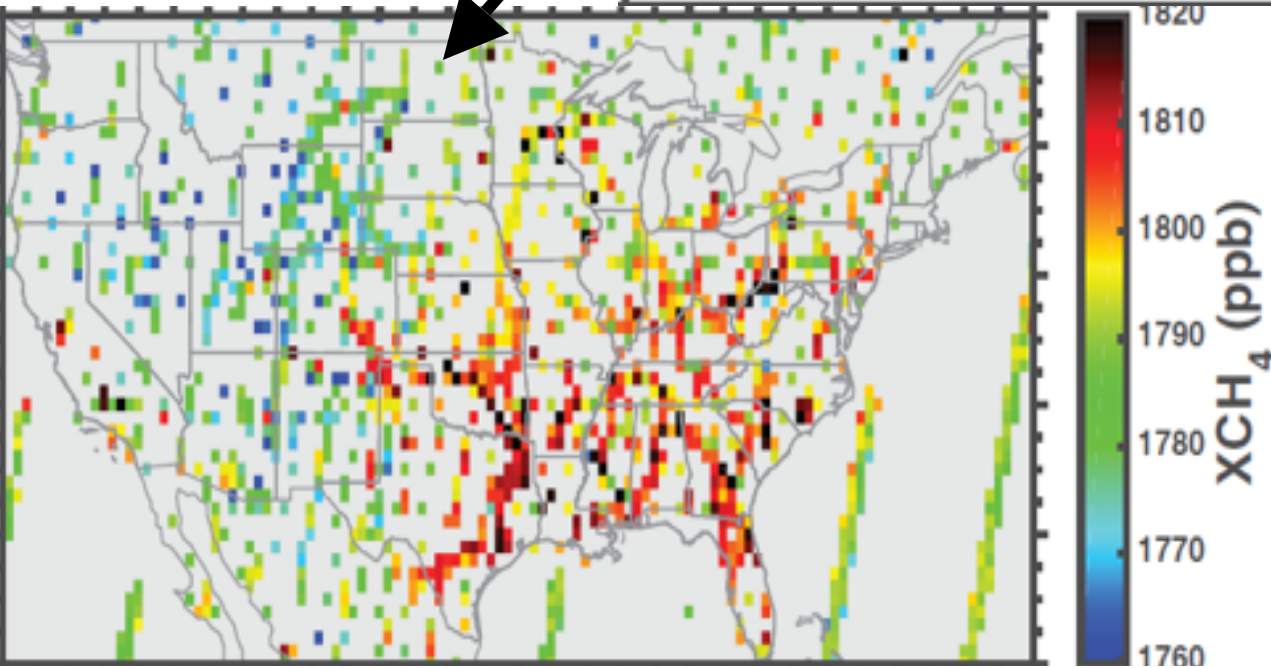
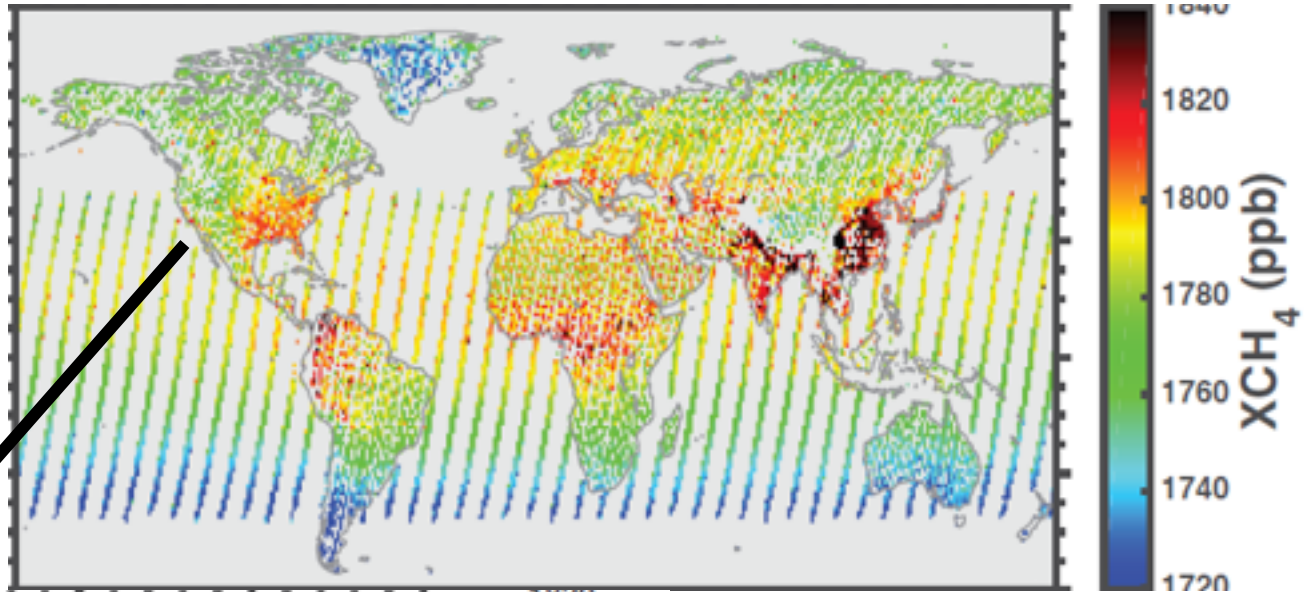


geoCARB (NASA), in formulation



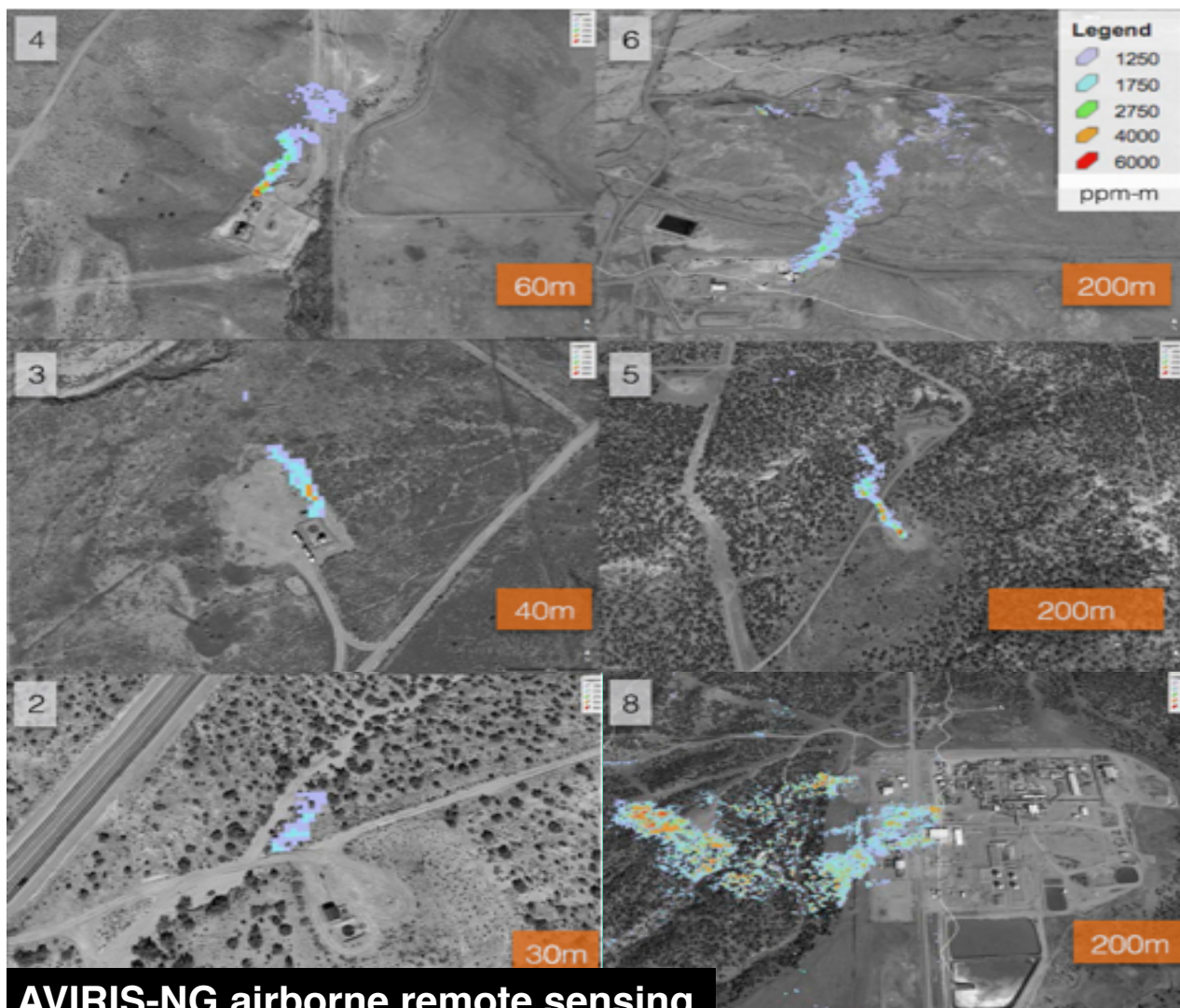
GOSAT satellite methane observations (2010-2013)

$X_{CH_4} \equiv$ column mole fraction



- 10x10 km² pixels, separated by 250 km
- 0.7% precision

Challenge of observing methane point sources at the facility scale: they are many and small and variable



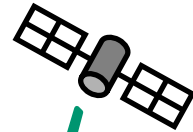
**AVIRIS-NG airborne remote sensing
of methane plumes in Four Corners
Frankenberg et al. [2016]**

GHGSat space-based observation of point sources

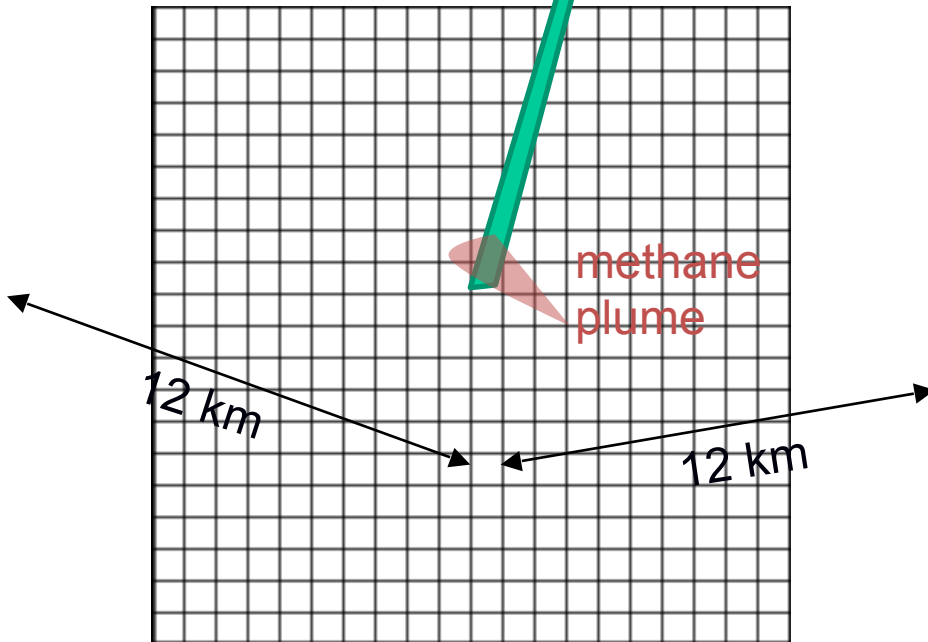
Effective pixel resolution of 50x50 m² over selected 12x12 km² scenes



First microsatellite
launched in June 2016



Smaller pixel resolution causes loss in precision...
but this is more than offset by lower dilution over pixel





Lom Pangar Dam, Cameroon

April 20th, 2017

GHGSat-D excess CH₄ column measurement



2014



Lom Pangar Dam, Cameroon

April 20th, 2017

GHGSat-D excess CH₄ column measurement



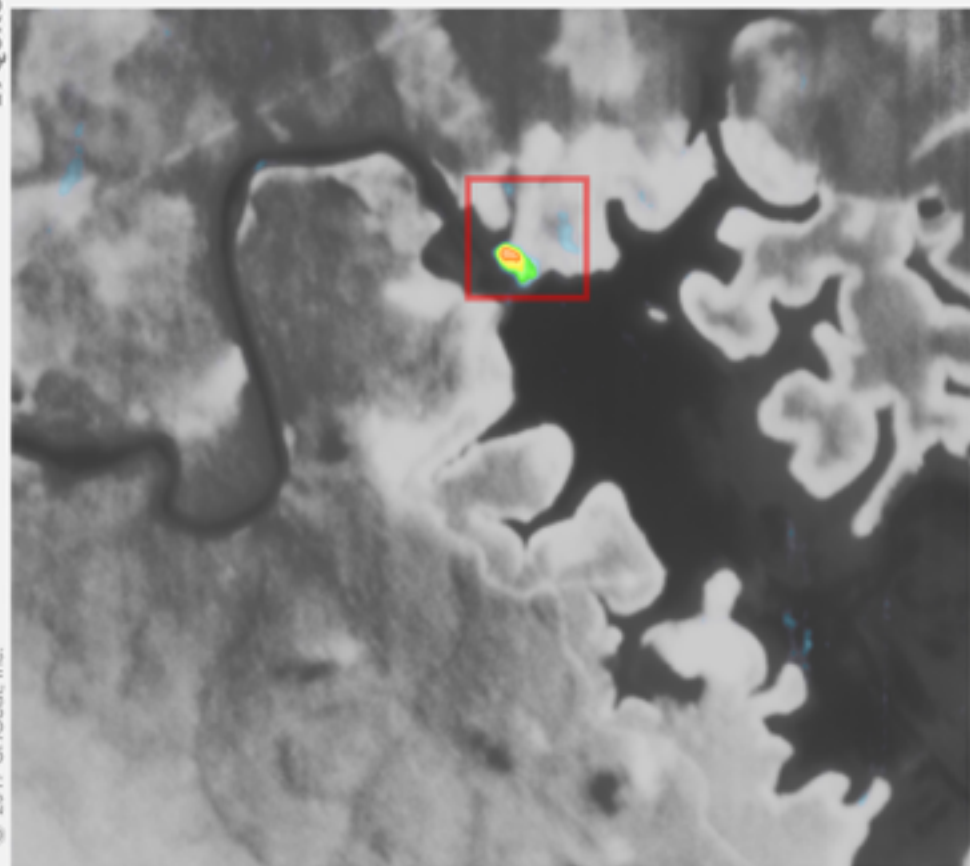
Lom Pangar Dam, Cameroon

April 20th, 2017

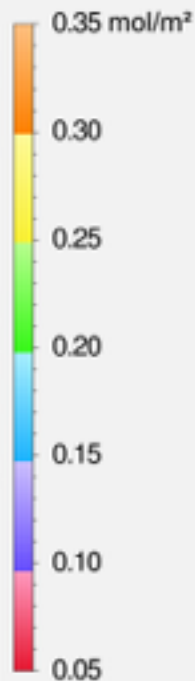
GHGSat-D excess CH₄ column measurement

17Q0RU1

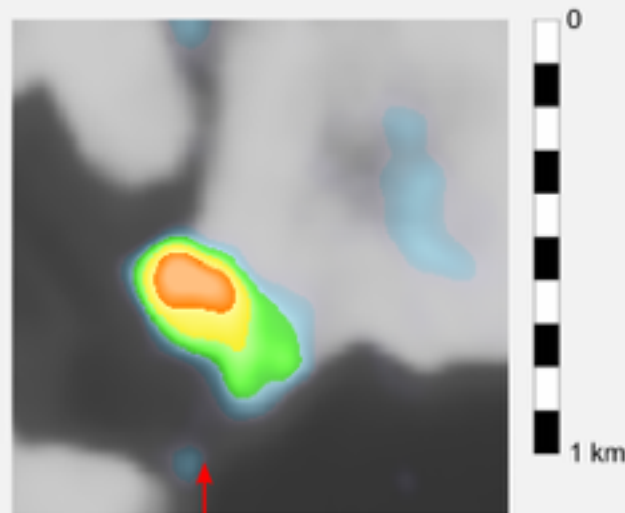
© 2017 GHGSat, Inc.



1 0 1 2 3 4 km



N
↑



Dam location
5°23'0.2"N 13°30'4.0"E

Background image: 1.6 μm reflectance
Timestamp: 2017-04-20 08:45:32 UTC