## Helpful Constants:

pptv = part per trillion  $(1 \times 10^{-12})$ ppbv = part per billion  $(1 \times 10^{-9})$ ppmv = part per million  $(1 \times 10^{-6})$ Surface area of a sphere:  $A = 4\pi r^2$  $1 \text{ m}^3 = 1000 \text{ l}$  $1 \text{ hPa} = 100 \text{ Pa} = 100 \text{ N/m}^2$ Gravitational acceleration  $g = 9.80 \text{ m/sec}^2$ Radius of the earth R = 6400 kmMean molecular mass of air  $M_a = 28.96$  g/mole Molecular mass of CO<sub>2</sub>:  $M_{CO_2} = 44$  g/mole  $N_2$  mixing ratio in the atmosphere  $C_{N_2} = 0.78$  mol/mol  $O_2$  mixing ratio in the atmosphere  $C_{O_2} = \text{oxygen} = 0.21 \text{ mol/mol}$ Dobson Unit (DU) =  $2.69 \times 10^{20}$  molec/m<sup>2</sup> Mass of the atmosphere:  $m_a = 4\pi R^2 P_s/g = 5.2 \times 10^{18}$  kg (surface pressure  $P_s$  in  $N/m^2$  = Pascals) Avogadros's number  $A_v = 6.02 \times 10^{23}$  (number of molecules in a mole) Ideal gas constant: R = 8.31 J/mol K

# Ideal Gas Law:

Mixing ratio:  $C_x = p_x/p = n_x/n_a$  (pressure, molar, or number density fraction) Moles of air in the atmosphere  $N_a = m_a/M_a$  ( $m_a$ : atmospheric mass, units consistent with  $M_a$ ) Moles of CO<sub>2</sub> atmosphere  $N_{CO_2} = C_{CO_2}N_a$  ( $N_a$  number of moles in the atmosphere)

- Mass of CO<sub>2</sub> in the atmosphere:  $m_{CO_2} = N_{CO_2} M_{CO_2}$
- Atmospheric column of x:  $\int n_x dz$  (molecules/area)

Ideal gas law: PV = NRT (Pressure P Pa (N/m<sup>2</sup>), Volume V (m<sup>3</sup>), moles N, temperature T (K) ) Number density of air :  $[M] = n_a = A_v N/V = A_v P/RT$  (in molec/m<sup>3</sup>, V in m<sup>3</sup>, N is moles,  $A_v$ : Avogadros's number).

Number density of air :  $n_a = P/(1.38 \times 10^{-19}T)$  ( $n_a$  in molec/ $cm^3$ , P in hPa and T in K). Species number density:  $n_x = C_x A_v P/RT = C_x n_a$ .

Saturated water vapor pressure  $P_{h_{20},sat}(T) = 6.112 \times e^{22.49 - (6142/T)}$  (in hPa, T in K).

 $RH(\%) = 100 \times P_{h_2o}/P_{h_2o,sat(T)}$ 

## Simple Models:

 $dm/dt = F_{in} + E + P - F_{out} - D - L$ 

If sources constant and all losses first order, can write: dm/dt = S - km, which has the solution,  $m(t) = m(0)e^{-kt} + \frac{S}{k}(1 - e^{-kt}).$ 

If loss of mass m is linear (first order): L = km.

Lifetime with respect to some first order loss process:  $\tau = \text{mass}/(\text{loss rate}) = 1/k$ .

Mass per unit area between pressure levels  $p_1$  and  $p_2$ ,  $M = (p_2 - p_1)/g$  ( $p_1$  and  $p_2$  in Pa)

Puff model of a pollution plume:  $d[X]/dt = E + P - L - D - k_{dil}([X] - [X]_b)$  where  $k_{dil}$  is a dilution rate constant, and  $[X]_b$  is the concentration of X in the background air.

#### **Biogeochemical cycles:**

Photosynthesis:  $CO_2 + H_2O + h\nu \rightarrow CH_2O + O_2$ , where CH<sub>2</sub>O represents biomass material. Burning, respiration, or decay:  $CH_2O + O_2 \rightarrow CO_2 + H_2O$ Total Dissolved Inorganic Carbon:  $[CO_2(aq)] = [CO_2 \cdot H_2O] + [HCO_3^-] + [CO_3^{2-}]$ Henry's Law: [X] is the molar concentration in water and  $P_X$  the pressure:  $[X] = K_H P_X$ .  $\begin{array}{l} \mathrm{CO}_2 + \mathrm{H}_2\mathrm{O} \Leftrightarrow \mathrm{CO}_2 \cdot \mathrm{H}_2\mathrm{O} \; (K_H = 0.03M/atm) \\ \mathrm{CO}_2 \cdot \mathrm{H}_2\mathrm{O} \Leftrightarrow \mathrm{HCO}_3^- + \mathrm{H}^+ \; (K_1 = 9 \times 10^{-7}M) \\ \mathrm{HCO}_3^- \Leftrightarrow \mathrm{CO}_3^{2-} + \mathrm{H}^+ \; (K_2 = 7 \times 10^{-10}M) \end{array}$ 

Molar fraction F of CO<sub>2</sub> that is in the atmosphere:  $F = N_{CO_2(g)}/(N_{CO_2(g)} + N_{CO_2(aq)})$ , where  $N_{CO_2(g)}$  is the number of moles in the atmosphere and  $N_{CO_2(aq)}$  the number of moles of inorganic carbon in the ocean.

Number of moles dissolved inorganic carbon in the ocean:  $N_{CO_2(aq)} = V_{oc}[CO_2(aq)]$ 

### **Radiative Transfer:**

optical depth:  $\delta = n\sigma L = N\sigma$  Here N is the "column" in molecules/area, n the concentration in molecules/volume, and  $\sigma$  the absorption coefficient of the molecule (or more generally, any absorber or scatterer, such as an aersosol).

Fractional Transmission  $T = e^{-\delta/\cos\theta}$  (where  $\delta$  is the optical depth in the vertical, and  $\theta$  the angle from the vertical.

$$\begin{split} F(x+dx) &= F(x) - dF_{abs} - dF_{scat} \\ \mathrm{dF}_{abs}: \text{ absorbed flux, the photon goes into the internal energy of the absorber} \\ \mathrm{dF}_{scat}: \text{ scattered flux, the photon bounces off in the forward or backward direction} \\ dF_{abs} &= n\sigma_{abs}F(x)dx \\ dF_{scat} &= n\sigma_{scat}F(x)dx \\ dF &= F(x+dx) - F(x) = -n(\sigma_{scat} + \sigma_{scat})F(x)dx \\ F(L) &= F(0)e^{-n(\sigma_{abs} + \sigma_{scat})L} \\ \delta &= ln\frac{F(0)}{F(L)} \\ \text{photolysis rate: } k = \int_{\lambda} q_x(\lambda)\sigma_x(\lambda)I_{\lambda}d\lambda, \text{ where } I_{\lambda} \text{ is the actinic flux as a function of wavelength.} \end{split}$$

#### Chemical Kinetics:

Standard bimolecular reaction:  $X + Y \rightarrow - > P + Q$ , d[X]/dt = -k[X][Y] where  $[X] = n_x$  and k is the rate constant. Standard Arrhenius form of reaction constant:  $k = k_0 e^{-E_a/RT}$ .  $E_a$  is the activation energy; R the gas constant; T the temperature in Kelvin.

Pressure dependent reaction:  $A + B + M \rightarrow AB + M$ Low Pressure Limit:  $\frac{d[AB]}{dt} = \frac{k_3k_5}{k_4}[A][B][M] = k_o[A][B][M]$ High Pressure Limit:  $\frac{d[AB]}{dt} = k_3[A][B] = k_{\infty}[A][B]$ Both Limits:  $\frac{d[AB]}{dt} = \frac{k_o[A][B][M]}{1 + \frac{k_o}{k_{\infty}}[M]}$